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# Photophysiological responses of marine diatoms to elevated CO<sub>2</sub> and decreased pH: a review

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**Abstract.** Diatoms dominate nearly half of current oceanic productivity, so their responses to ocean acidification are of general concern regarding future oceanic carbon sequestration. Community, mesocosm and laboratory studies show a range of diatom growth and photophysiological responses to increasing  $pCO_2$ . Nearly 20 studies on effects of elevated  $pCO_2$  on diatoms have shown stimulations, no effects or inhibitions of growth rates. These differential responses could result from differences in experimental setups, cell densities, levels of light and temperature, but also from taxon-specific physiology. Generally, ocean acidification treatments of lowered pH with elevated  $CO_2$  stimulate diatom growth under low to moderate levels of light, but lead to growth inhibition when combined with excess light. Additionally, diatom cell sizes and their covarying metabolic rates can influence responses to increasing  $pCO_2$  and decreasing pH, although cell size effects are confounded with taxonomic specificities in cell structures and metabolism. Here we summarise known diatom growth and photophysiological responses to increasing  $pCO_2$  and decreasing pH, and discuss some reasons for the diverse responses observed across studies.

Additional keywords: Bacillariophyceae, CO<sub>2</sub>, diatom, ocean acidification, photoinhibition, photosynthesis.

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# Introduction

Increasing atmospheric  $CO_2$  due to anthropogenic activities affects terrestrial photosynthesis, but is also causing p $CO_2$  to rise and pH to drop in the surface oceans, which influence marine primary producers (Beardall *et al.* 2009; Riebesell and Tortell 2011; Gao *et al.* 2012*a*), in a yet more complicated way due to the concurrent changes in seawater chemistry and ocean mixing.

Diatoms are suggested to have evolved between 100 and 200 million years ago (Sims et al. 2006), when atmospheric  $CO_2$  is thought to have been ~2000 ppmv (Veron 2008), compared with current levels of ~400 ppmv and projected endcentury levels of 750-1000 ppmv. The extant diatoms now contribute ~20% of the organic carbon generated globally each year by photosynthesis (Field et al. 1998). They exist both as phytoplankton and as benthic algae. At least some diatom species operate metabolic pathways unusual among studied phytoplankton, including a urea cycle (Allen et al. 2011) and a C<sub>4</sub> carboxylation path (Reinfelder et al. 2004; Haimovich-Dayan et al. 2013). Diatoms run highly efficient CO<sub>2</sub> concentrating mechanisms (CCMs) to achieve a high ratio of carboxylation to oxygenation (Raven et al. 2011; Reinfelder 2011). They tolerate high levels of UV radiation (Guan and Gao 2008; Wu et al. 2012a), enjoy a low susceptibility to photoinactivation of PSII compared with other phytoplankters (Six et al. 2007, 2009; Key et al. 2010; Wu et al. 2011) and successfully exploit variable light (Lavaud et al. 2004,

2007). They are, as a net result, by far the most successful group of eukaryotic aquatic primary producers, not only in terms of primary production but also in their number of species and their capacities to acclimate to environmental changes with diversified metabolisms. Diatom growth rates correlate closely with their size, decreasing almost linearly with the log of increasing cell volume, which ranges across eight orders of magnitude, with cell diameters ranging from ~2  $\mu$ m to a few mm (Finkel *et al.* 2010). Diatoms have responded to past climate change through successions of differently sized cells, with a trend towards smaller cells under higher temperatures over the past 65 million years (Falkowski and Oliver 2007; Finkel *et al.* 2007).

The ongoing ocean acidification triggered by increasing atmospheric  $CO_2$  concentration alters seawater carbonate chemistry, the availabilities and toxicities of nutrients (Millero *et al.* 2009). Ocean warming will in concert tend to drive increased stratification, decreasing the thickness of the upper mixing layer and lowering transport of nutrients from interior or deeper layers to the surface ocean (Doney 2006; Steinacher *et al.* 2010). These changes will differentially affect differently sized diatom species (Flynn *et al.* 2012), and thereby alter sinking rates and organic carbon export (Finkel *et al.* 2010). Therefore, growth and physiological responses of diatoms to elevated  $CO_2$  concentrations have gained attention (Riebesell *et al.* 1993; Burkhardt and Riebesell 1997; Burkhardt *et al.* 1999). Stimulative, neutral and inhibitory effects of elevated  $CO_2$  Burkhardt *et al.* 1999; Chen and Gao 2003; Kim *et al.* 2006; Wu *et al.* 2010; Yang and Gao 2012; Li and Campbell 2013) (Table 1). In this review, we focus on the effects of elevated  $CO_2$  and lower pH on diatom photophysiology, by summarising studies of their growth and photophysiological responses to elevated  $CO_2$  under different experimental conditions, and discuss some potential mechanisms to resolve the diversity of responses observed across studies to date.

# **Ocean acidification**

The oceans are presently absorbing ~25 million tons of  $CO_2$  from the atmosphere each day, an important role in counteracting global warming (Sabine *et al.* 2004). This dissolution of  $CO_2$ from the air is, however, acidifying the oceans (Doney 2006; Gattuso and Hansson 2011).

The exchange of CO<sub>2</sub> between the sea and atmosphere depends on temperature, salinity, physical mixing of seawater, respiration and photosynthesis. Therefore, CO<sub>2</sub> fluxes change horizontally due to physical, chemical and biological properties of waters. When CO<sub>2</sub> dissolves in seawater, it combines with water to form carbonic acid  $(CO_2 + H_2O \rightarrow H_2CO_3)$ which dissociates to bicarbonate  $(H_2CO_3 \rightarrow H^+ + HCO_3^-)$ , discharging protons (H<sup>+</sup>) and so ultimately attaining a state of new equilibrium. However, as the  $H^+$  concentration increases with CO<sub>2</sub> dissolution, the H<sup>+</sup> releases can partially reverse the secondary dissociation reaction, leading to a decrease in carbonate ions  $(H^+ + CO_3^{2-} \rightarrow HCO_3^-)$ . Typical changes linked with ocean acidification are therefore increased concentrations of pCO<sub>2</sub>, H<sup>+</sup> and HCO<sub>3</sub><sup>-</sup>, decreases in the concentration of  $CO_3^{2-}$  and decreases in the CaCO<sub>3</sub> saturation state (Gattuso et al. 2010). Since the beginning of the industrial revolution, the pH of oceanic surface seawater has already dropped by ~0.1 unit due to atmospheric CO<sub>2</sub> rise (Caldeira and Wickett 2003),

#### Table 1. Effects of elevated CO<sub>2</sub> concentrations reported in diatoms

Effects	Authors	Species	Aspects (% change)
Stimulative	(Kim et al. 2006)	Skeletonema costatum	Growth (50%)
	(Wu et al. 2010)	Phaeodactylum tricornutum	Growth (5.2%)
	(King et al. 2011)	Attheya sp.	Growth (31%)
	(Low-Décarie et al. 2011)	Navicula pelliculosa	Growth (100%)
	(Sun et al. 2011)	Pseudo-nitzschia multiseries	Photosynthesis (21%)
	(Gao <i>et al.</i> 2012 <i>b</i> )	Skeletonema costatum	Growth (12%)
	(Gao <i>et al.</i> 2012 <i>b</i> )	Phaeodactylum tricornutum	Growth (16%)
	(Yang and Gao 2012)	Thalassiosira pseudonana	Growth (6.5%)
		Thalassiosira pseudonana	Photosynthesis (17%)
	(McCarthy et al. 2012)	Thalassiosira pseudonana CCMP 1014	Growth (50%); PSII electron transport under low light (18%)
	(McCarthy et al. 2012)	Thalassiosira pseudonana CCMP 1335	Growth (14%); PSII electron transport under low light (72%)
	(Li and Campbell 2013)	Thalassiosira pseudonana CCMP 1335	Growth under low to optimal light (25%)
Unaffected	(Riebesell et al. 1993)	Ditylum brightwellii	Growth
		Thalassiosira punctigera	Growth
		Rhizosolenia cf. alata	Growth
	(Chen and Gao 2004a)	Skeletonema costatum	Growth; photosynthesis
	(Kim et al. 2006)	Nitzschia spp.	Growth
	(Crawfurd et al. 2011)	Thalassiosira pseudonana CCMP 1335	Growth
	(Boelen et al. 2011)	Chaetoceros brevis	Growth
	(Gao <i>et al.</i> 2012 <i>b</i> )	Diatoms <sup>A</sup>	Growth (moderate light)
	(Yang and Gao 2012)	Thalassiosira pseudonana	Growth; photoinhibition
	(Ihnken et al. 2011)	Chaetoceros muelleri	Growth (moderate light)
	(Boelen et al. 2011)	Chaetoceros brevis	Growth; photosynthesis
Negative	(Wu et al. 2010)	Phaeodactylum tricornutum	Photoinhibition (14.5%); dark respiration (33.7%)
	(Low-Décarie et al. 2011)	Nitzschiapalea	Growth (-67%)
	(Gao <i>et al.</i> 2012 <i>b</i> )	Skeletonema costatum	Growth (high light; -12%);
	(Gao <i>et al.</i> 2012 <i>b</i> )	Phaeodactylum tricornutum	Growth (highlight; -10%); photorespiration (26%)
		Thalassiosira pseudonana	Growth (high light; -16%); photorespiration (24%)
	(Torstensson et al. 2012)	Navicula directa	Growth (-5%)
	(Yang and Gao 2012)	Thalassiosira pseudonana	Dark respiration (-35%)
	(Ihnken et al. 2011)	Chaetoceros muelleri	Growth (-10%)
	(Li and Campbell 2013)	Thalassiosira pseudonana CCMP 1335	Growth under excess light $(-13\%)$
	(Sobrino et al. 2008)	Thalassiosira pseudonana	Photoinactivation (25%)
	(Mejía et al. 2013)	Thalassiosira pseudonana	Decreased silicification (12%)
	(Mejía <i>et al.</i> 2013)	Thalassiosira weissflogii	Decreased silicification (90%)

<sup>A</sup>Phaeodactylum tricornutum, Skeletonema costatum, Thalassiosira pseudonana.

equivalent to about a 30% increase in the H<sup>+</sup> concentration. With a further increase of CO<sub>2</sub> concentration in the atmosphere to 800-1000 ppmv under the IPCC A1F1 scenario (Houghton et al. 2001), by the end of this century, pH of the surface oceans will decrease by another 0.3-0.4 units (Feely et al. 2004; Sabine et al. 2004; Orr et al. 2005), thus, increasing [H<sup>+</sup>] by 100–150%. Consequently, organisms in the euphotic zone will be exposed to a higher CO<sub>2</sub> and a lower pH, and their physiologies will respond to changes in seawater carbonate chemistry, as well as to secondary changes in ionic speciation and cell surface chemistry driven by decreasing pH (Millero et al. 2009; Flynn et al. 2012; Hervé et al. 2012; Sugie and Yoshimura 2013). These chemical changes can directly affect physiology of marine organisms (Pörtner and Farrell 2008), but can also indirectly influence organismal responses to other environmental factors including UV radiation (Sobrino et al. 2008; Gao et al. 2009; Chen and Gao 2011; Li et al. 2012a), light (Bartual and Galvez 2002; Sobrino et al. 2008; McCarthy et al. 2012; Li and Campbell 2013), temperature change (Pörtner and Farrell 2008; Zou et al. 2011) or nutrients (Burkhardt and Riebesell 1997; Burkhardt et al. 1999; Riebesell and Tortell 2011; Li et al. 2012b).

## Growth responses

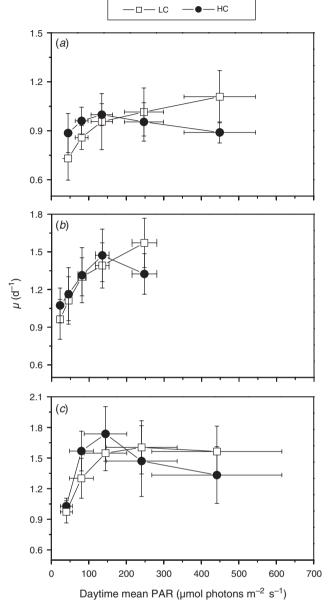
Dissolved inorganic carbon (DIC) in surface seawater, at present, is ~100–200 times that of  $CO_2$  in the atmosphere, but most seawater DIC is HCO<sub>3</sub>, with CO<sub>2</sub> typically accounting for less than 1% in pelagic waters (Gattuso et al. 2010). In addition, CO2 in seawater diffuses ~8000 times slower than in air, which can kinetically limit marine photosynthetic carbon fixation (Raven 1993; Riebesell et al. 1993; Morel et al. 1994). Growth of diatom species can, in turn, be limited by the availability of CO<sub>2</sub> (Riebesell et al. 1993), and oceanic primary production might thus be enhanced by increasing atmospheric CO<sub>2</sub> concentration (Hein and Sand-Jensen 1997; Schippers et al. 2004; Riebesell and Tortell 2011). However, the growth rate of diatom-dominated phytoplankton assemblages was not affected by an elevated pCO<sub>2</sub> concentration of 800 µatm during 2-5 days shipboard incubation under ~30% of incident sunlight (Tortell et al. 2000). Growth of Skeletonema costatum was not stimulated by an enriched CO<sub>2</sub> concentration (800 µatm) under laboratory conditions (Burkhardt and Riebesell 1997; Chen and Gao 2003, 2004a), but was enhanced in a mesocosm at an elevated CO<sub>2</sub> concentration of 750 µatm (Kim et al. 2006). Growth of the diatoms Phaeodactylum tricornutum (Schippers et al. 2004; Wu et al. 2010), Navicula pelliculosa (Low-Décarie et al. 2011) and Attheya sp. (King et al. 2011) were also enhanced under elevated CO<sub>2</sub> levels under laboratory conditions. However, in the diatom Chaetoceros muelleri, low-light treatments showed lower growth rates under elevated CO<sub>2</sub> conditions, but no CO<sub>2</sub> or pH effect was recorded under high light exposure (Ihnken et al. 2011). Under similar laboratory conditions, while growth of Thalassiosira pseudonana (CCMP 1335) was not stimulated at the elevated CO<sub>2</sub> levels of 760 (Crawfurd et al. 2011) or 1000 µatm (Yang and Gao 2012), T. pseudonana (CCMP 1014 and 1335) grew faster under low (McCarthy et al. 2012) to moderate light (Li and Campbell 2013) with pCO<sub>2</sub> of 750 µatm,

but under higher light T. pseudonana CCMP 1335 suffered growth inhibition (Li and Campbell 2013). Growth rates of the diatoms S. costatum (CCMA110), P. tricornutum (CCMA 106) and T. pseudonana (CCMP 1335), when grown under different levels of sunlight and elevated CO<sub>2</sub> of 1000 µatm, were stimulated under lower light levels (5-30% surface daytime mean solar PAR), but inhibited under higher light levels, with a drop in the PAR threshold for growth saturation (Fig. 1). These results show that elevated CO<sub>2</sub> and light levels interact to affect diatom growth responses to ocean acidification, which may explain the different results obtained under different experimental setups, at least for the same species (Table 1). Some of these growth responses may relate to CO<sub>2</sub> dependent changes in the cellular susceptibility to photoinactivation of PSII (Li and Campbell 2013) (see below). Since future shoaling of upper-mixed-layer (UML) depths is expected to expose phytoplankton to increased solar irradiance, marine primary producers within UML are expected to suffer from enhanced light stress. However, both low-light CO2 growth enhancement and high-light CO2 growth inhibition could occur even within a single daytime solar cycle or a vertical mixing path, making the net community outcomes difficult to predict. Boelen et al. (2011) did not find any interactive effects of elevated CO<sub>2</sub> concentration and changing light levels nor fluctuating light on the growth and photosynthesis in the Antarctic diatom Chaetoceros brevis. Conversely, frequencies of light fluctuation that mimic different mixing regimes affect a coccolithophore's response to ocean acidification (Jin *et al.* 2013a), implying an interactive effect of light fluctuation and ocean acidification, that could impose an additional layer of influence on the net species and community responses to increasing pCO<sub>2</sub>.

## Inorganic carbon acquisition mediated by elevated CO<sub>2</sub>

Diatom Rubisco shows comparatively high CO<sub>2</sub> affinity and  $CO_2/O_2$  selectivity, and is served by  $CO_2$  concentrating mechanisms (CCMs) to supply CO2 to Rubisco, and thereby diminish photorespiration (Roberts et al. 2007a). CCMs differ among studied diatom species. Thalassionema nitzschioides (Trimborn et al. 2009), Thalassiosira weissflogii and P. tricornutum (Burkhardt et al. 2001) actively take up both CO<sub>2</sub> and HCO<sub>3</sub><sup>-</sup>, whereas *Thalassiosira punctigera* exclusively uses free CO<sub>2</sub> (Elzenga et al. 2000). T. pseudonana, though lacking periplasmic (known also as extracellular) carbonic anhydrase (eCA), can take up HCO<sub>3</sub><sup>-</sup> directly (Nimer et al. 1997; Elzenga et al. 2000; Nakajima et al. 2013), and indeed uses HCO<sub>3</sub><sup>-</sup> as the dominant substrate for photosynthesis even under increased pCO<sub>2</sub> (Hopkinson et al. 2013; Isensee et al. 2013). The eCA and intracellular carbonic anhydrase (iCA) facilitate Ci acquisition or utilisation by catalysing the interconversion of  $CO_2$  and  $HCO_3^-$ .

The activity of eCA can be downregulated under elevated  $CO_2$  concentrations relevant to climate change (Burkhardt *et al.* 2001; Chen and Gao 2003; Rost *et al.* 2003; Crawfurd *et al.* 2011). Therefore, active transport or use of  $HCO_3^-$  could be lowered under elevated  $CO_2$ . Different growth conditions can therefore bring about different efficiencies of algal CCMs or changing preferences for  $CO_2$  or  $HCO_3^-$  (Korb *et al.* 1997; Nimer *et al.* 



**Fig. 1.** Light-dependence of diatom growth responses to elevated CO<sub>2</sub> (HC, 1000 µatm, pH<sub>T</sub> 7.68) compared with ambient CO<sub>2</sub> level (LC, 390 µatm, pH<sub>T</sub> 8.02). Growth rates were stimulated by elevated CO<sub>2</sub> under low to moderate PAR, but inhibited under higher PAR levels. The PAR thresholds for the transition from CO<sub>2</sub> growth stimulation, under lower light, to CO<sub>2</sub> growth inhibition, under higher light, were ~160 for *Phaeodactylum tricornutum* (*a*), 125 for *Thalassiosira pseudonana* (*b*) and 178 amol photons m<sup>-2</sup> s<sup>-1</sup>, for *Skeletonema costatum* (*c*). These threshold light levels for the transition for CO<sub>2</sub> growth stimulation correspond to 22–36% of the incident surface solar PAR levels and are equivalent to PAR levels at 26–39 m depth in the South China Sea. The semi-continuous cultures were maintained under the sun and diluted every 24 h to ensure stability of cell density ranges and the seawater carbonate system (from Gao *et al.* 2012*b*).

1997; Burkhardt *et al.* 2001). In microalgae and in cyanobacteria, very high  $CO_2$  concentrations of up to 50000 µatm (or ppmv in air) turn off CCMs (Kaplan *et al.* 1980; Tsuzuki and Miyachi

1989; Raven 1991; Matsuda *et al.* 2001). CCM induction is closely related to the intracellular Ci pool, ambient CO<sub>2</sub> levels and oxygen availability (Woodger *et al.* 2005). CO<sub>2</sub> levels (up to 1000  $\mu$ atm) relevant to future CO<sub>2</sub> levels projected for 2100 have been confirmed to partially downregulate CCMs in marine diatoms (Chen and Gao 2003; Trimborn *et al.* 2009; Wu *et al.* 2010, 2012*a*; Yang and Gao 2012) and to lower reliance upon an intracellular labile carbon pool (Isensee *et al.* 2013). Downregulation of CCMs can include decreased CO<sub>2</sub> affinity resulting in an increased requirement for pCO<sub>2</sub> to support photosynthesis, inhibition of carbonic anhydrase activity, depressed HCO<sub>3</sub><sup>--</sup> transport, and downregulation of PEPCase and PEPCKase (Reinfelder *et al.* 2000; Giordano *et al.* 2005; Roberts *et al.* 2007*a*, 2007*b*; Raven 2010; Reinfelder 2011). Such CCM downregulation was found to be synchronised with diurnal

Gao 2004*b*). CCMs in diatoms might connect to multiple metabolic pathways which differ among diatom species of differing physiology or sizes. In *P. tricornutum*, cAMP metabolism is involved to control CCM under elevated CO<sub>2</sub> levels (Harada *et al.* 2006). *T. weissflogii* appears to run a C<sub>3</sub>–C<sub>4</sub>-intermediate photosynthesis (Roberts *et al.* 2007*a*), which may concentrate Ci through incorporation into an organic C<sub>4</sub> carbon compound, before Rubisco-aided carboxylation (Reinfelder *et al.* 2000, 2004). However, in *P. tricornutum*, the C<sub>4</sub> path way was recently suggested to contribute to pH homeostasis or excitation dissipation, but not to a CCM function (Haimovich-Dayan *et al.* 2013).

photosynthetic performance in the diatom S. costatum (Chen and

CCMs consume energy (Raven 1991; Bouma et al. 1994; Crawfurd et al. 2011; Hopkinson et al. 2011). Active uptake of  $HCO_3^-$  and  $CO_2$  is supported by cyclic and linear electron transport in cyanobacteria (Li and Canvin 1998). Pseudocyclic electron flow through the Mehler reaction can also contribute (Sültemeyer et al. 1993). The energisation mechanisms of diatom CCMs are as yet unclear. It addition to the initial uptake, it takes further energy to maintain high intracellular Ci levels by counteracting CO<sub>2</sub> efflux (Sukenik et al. 1997; Tchernov et al. 1997) although in diatoms tested to date efflux rates appear small (Burkhardt et al. 2001; Trimborn et al. 2009). The major energy expenditure by the CCMs in diatoms is active Ci transport, and a doubling of ambient [CO2] could save ~20% of the CCM-related energy expenditure in several diatom species (Hopkinson et al. 2011). Under elevated  $CO_2$  concentrations, the growth enhancement under limiting light levels could be partially due to downregulation of CCMs, thereby lowering energy costs (Raven and Johnston 1991; Gao et al. 2012b). Alternatively, since CCMs can also be downregulated under low light, elevated pCO<sub>2</sub> could have stimulated the low light growth of diatoms due to both increased availability of CO<sub>2</sub> and savings on the energy cost of CCM operation.

# Photosynthetic responses

Light energy captured and delivered via photochemical processes powers the active transport of  $CO_2$  and  $HCO_3$  in cyanobacteria and microalgae (Sültemeyer *et al.* 1993; Sukenik *et al.* 1997; Li and Canvin 1998), and then assimilatory carboxylation. Elevated  $CO_2$  concentration had no significant effect upon pigment contents nor upon the effective absorbance cross-section serving photosystem II photochemistry ( $\sigma_{PSII}$ ) in *S. costatum*, *P. tricornutum* or *T. pseudonana* (Chen and Gao 2004*a*; Wu *et al.* 2010; Crawfurd *et al.* 2011; McCarthy *et al.* 2012; Li and Campbell 2013). Furthermore, elevated pCO<sub>2</sub> had only limited effects on levels of the major protein complexes mediating photosynthesis across multiple species of centric diatoms, grown under low to saturating light (McCarthy *et al.* 2012; Li and Campbell 2013).

The photochemical quantum yield of S. costatum, P. tricornutum and T. pseudonana decreases faster with increasing levels of PAR under elevated, than under ambient CO<sub>2</sub> levels (Gao et al. 2012b). Non-photochemical quenching (NPO), however, increases faster in the high-CO<sub>2</sub> grown cells with increasing light levels compared with the ambient CO<sub>2</sub> grown cells (Gao et al. 2012b). Modulating NPQ helps diatoms, like other photoautoauphs, to withstand high or fluctuating levels of PAR (Niyogi et al. 2005; Lavaud et al. 2007; Zhu and Green 2010; Wu et al. 2012b). T. pseudonana employs NPQ to cope with light stress, even under elevated CO<sub>2</sub> levels, more effectively than does a strain of P. tricornutum (Yang and Gao 2012), so that photoinhibition of electron transport was observed in P. tricornutum, but not T. pseudonana, when grown under elevated  $CO_2$  of 1000 µatm (Wu et al. 2010). These differential responses between two model diatoms show taxon-specific mechanisms in coping with the combined impacts of ocean acidification and light stress.

At a functional level, the diatoms P. tricornutum and T. pseudonana grown under elevated CO<sub>2</sub> of 1000 µatm, at subsaturating photosynthetically active radiation, showed an increase in photosynthetic carbon fixation rate per cell of more than 20% (Wu et al. 2010; Yang and Gao 2012). Their growth rate was, however, only enhanced by ~5% in P. tricornutum and was unaffected in T. pseudonana. Enhanced respiratory and photorespiratory carbon losses under elevated CO<sub>2</sub> are likely responsible for this discrepancy (Wu et al. 2010; Gao et al. 2012b; Yang and Gao 2012). In the toxic diatom Pseudo-nitzschia multiseries, maximum carbon fixation rates per cell also increased with elevated CO<sub>2</sub> levels, although the apparent light use efficiency was not affected (Sun et al. 2011). In Cylindrotheca closterium f. minutissima, when grown at 1000 µatm CO<sub>2</sub> under sunlight, rates of electron transport and O<sub>2</sub> evolution dropped compared with the cells grown at the ambient CO<sub>2</sub> concentration (Wu et al. 2012a).

### PSII photoinactivation and UV responses

Diatoms, like all photoautotrophs, suffer light- and UVdependent photoinactivation of their PSII centres (Kok 1956). To maintain their photosynthesis in the face of light-dependent photoinactivation, diatoms must use a metabolically expensive PSII repair cycle (Aro *et al.* 1993) to proteolytically remove photoinactivated protein subunits (Nixon *et al.* 2010; Nagao *et al.* 2012; Campbell *et al.* 2013) and replace them with newly synthesised subunits (Edelman and Mattoo 2008). In comparison with other phytoplankton groups, diatoms enjoy a relatively low susceptibility to photoinactivation of their PSII (Key *et al.* 2010; Wu *et al.* 2011, 2012*b*). In *T. pseudonana* CCMP 1335, however, the primary susceptibility to photoinactivation of PSII changes under elevated pCO<sub>2</sub> (Sobrino et al. 2008; McCarthy et al. 2012; Li and Campbell 2013). As a net result, cells under high pCO<sub>2</sub> and high light incur an increased metabolic expense to accelerate PSII protein cycling, to counter increased photoinactivation (G Li, DA Campbell, unpubl. data). This increased metabolic cost to maintain PSII function is a possible explanation for the pattern of growth stimulation under elevated pCO<sub>2</sub> under low to moderate light, but growth inhibition under excess light (Gao et al. 2012b; McCarthy et al. 2012; Li and Campbell 2013). We are as yet unsure as to the mechanism(s) for the changes in susceptibility to primary photoinactivation under elevated pCO<sub>2</sub>. Decreased silification under elevated pCO<sub>2</sub> (Mejía et al. 2013) might alter cellular optics. Or, a drop in excitation dissipation capacity, as reported in cyanobacteria (Tchernov et al. 1997) and now suggested in diatoms (Haimovich-Dayan et al. 2013) could result if the CCM is partly downregulated. Evidence in this direction is that the content of the reactive-oxygen toxicity indicator malondialdehyde increases in T. pseudonana CCMP 1335 growing under elevated pCO<sub>2</sub> (Li and Campbell 2013), consistent with a downregulation of paths with photoprotective roles under elevated pCO<sub>2</sub>.

Solar UV radiation (UVR, 280–400 nm) affects phytoplankton physiology and primary productivity (Häder 2011; and literatures cited therein). In *T. pseudonana*, acclimation to UVR, partially relieved the increased susceptibility to photoinhibition under elevated pCO<sub>2</sub> (Sobrino *et al.* 2008), consistent with a hormetic protective induction of reactive oxygen species (ROS) detoxification by UV acclimation. The effect of UV-B irradiance (280–320 nm) on *P. tricornutum* was counteracted under ocean acidification conditions (Li *et al.* 2012*a*). *Cylindrotheca closterium* f. *minutissima* did not show any significant growth response to solar UVR after acclimation to solar radiation, though a combination of UVR and elevated  $CO_2$  concentration led to significant drop in maximal electron transport (Wu *et al.* 2012*a*).

## **Respiratory responses**

Altered seawater carbonate chemistry due to ocean acidification could perturb energy requirements for the diatom cells, leading to changes in respiration. Mitochondrial respiration indeed increases under ocean acidification conditions of 1000  $\mu$ atm pCO<sub>2</sub> (pH 7.8) by ~34% in *P. tricornutum* (Wu *et al.* 2010) and by 35% in *T. pseudonana* (Yang and Gao 2012). Increased acidity of seawater associated with increased pCO<sub>2</sub> could disturb cell surface (Flynn *et al.* 2012) or even intracellular pH stability, so that phytoplankton cells may need to allocate additional energy to transport ions against the acid–base perturbation. Cell surface effects of increasing pCO<sub>2</sub> and decreasing pH will vary with cell size and with the co-varying cellular metabolic rate (Flynn *et al.* 2012). Thus, increasing pCO<sub>2</sub> is likely to increase the influences of cell size on phytoplankton responses to environmental forcings (Finkel *et al.* 2010; Flynn *et al.* 2012).

Photorespiration and electron flows to oxygen can be important in photoprotection and short-term responses to excess light in diatoms (Wingler *et al.* 2000; Waring *et al.* 2010). Both *P. tricornutum* and *T. pseudonana* showed enhanced photorespiration by up to 23-27% under elevated CO<sub>2</sub> or ocean acidification conditions (Gao *et al.* 2012*b*). Cells of *T. pseudonana* grown under the elevated CO<sub>2</sub> level of 1000 µatm showed higher carbon fixation rates but a lower net O<sub>2</sub> evolution rate compared with cells grown under the ambient CO<sub>2</sub>, although the cells exhibited equivalent electron transfer rates from PSII (Yang and Gao 2012), providing evidence for enhanced photorespiratory or pseudocyclic re-consumption of O<sub>2</sub> released from PSII, under ocean acidification. Enhanced excretion of organic compounds due to photorespiration can connect to production of transparent exopolymers in phytoplankton communities that include diatoms (Engel 2002). These metabolic pathways could result in discrepant effects of ocean acidification on different species or under different light levels.

## Effects on diatom communities across diverse habitats

Although photosynthesis of diatoms is likely to be stimulated by increased availability of CO<sub>2</sub>, lower pH might increase their respiration (Wu *et al.* 2010; Yang and Gao 2012) and their costs for photoprotection, therefore, the net effect of ocean acidification on diatoms will depend on multiple environmental forcings and possibly species-specific metabolic pathways.

Community level responses to rising pCO<sub>2</sub> and temperature vary across oceanic regions. In the north-east Atlantic and North Sea a 50 year (1960-2009) time series survey revealed a decline of dinoflagellate abundance, whereas diatoms showed relatively constant richness (Hinder et al. 2012). The transition over the half century was attributed to ocean warming and windy conditions. In contrast, elevated temperatures lowered the short-term abundance of diatoms in a North Atlantic Bloom incubation study (Feng et al. 2009), although CO<sub>2</sub> changes had no apparent effect. A shipboard incubation study that examined rising temperature and CO2 in two natural Bering Sea assemblages also found large community shifts away from diatoms towards nanoflagellates in the 'greenhouse' treatment (Hare et al. 2007), though diatom-dominated phytoplankton growth increased in the Ross Sea under elevated CO<sub>2</sub> levels (Tortell et al. 2008).

Across a CO<sub>2</sub> and pH gradient off the volcanic island of Vulcano (Mediterranean, NE Sicily), periphyton communities altered significantly as CO<sub>2</sub> concentrations increased, with significant increases in chlorophyll *a* concentrations and in diatom abundance (Johnson *et al.* 2013).

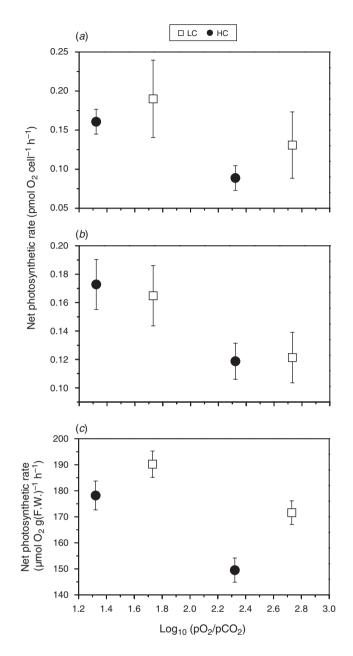
Feng *et al.* (2010) found no interactive effects of light and  $CO_2$  on community photosynthesis during an experiment using a Ross Sea diatom/*Phaeocystis* assemblage, but the diatom community structure shifted away from small pennate diatoms towards much larger centric diatoms. In the very different conditions of the South China Sea a shipboard incubation combining elevated  $CO_2$  concentration and near surface solar irradiances, showed decreased photosynthesis while diatom abundance declined (Gao *et al.* (2012*b*).

In hypoxic seawaters, algae may experience large changes in the ratio of  $pO_2$  to  $pCO_2$  or respiration index (RI =  $log_{10}$  ( $pO_2/$  $pCO_2$ )), which is predicted to decline in future oceans (Brewer and Peltzer 2009). Together with lower pH, hypoxic areas represent a future situation of combined ocean acidification and deoxygenation. There has been little documented on the interactive effects of these two climate-change factors on diatoms nor upon other marine primary producers. However, changes in RI could affect net photosynthesis (Fig. 2) (Gao *et al.* 2012*b*; Xu and Gao 2012). For diatoms grown under either elevated or ambient levels of CO<sub>2</sub>, net photosynthetic O<sub>2</sub> evolution decreases with increased RI (Fig. 2*a*, *b*).

# Future efforts

Based on the work summarised here, priorities for future study on diatoms include the following.

- Response of frustule mineralisation to elevated CO<sub>2</sub> concentrations. Biogenic silicate content of some diatoms decreases under ocean acidification conditions (Hervé *et al.* 2012; Tatters *et al.* 2012; Mejía *et al.* 2013). However, we do not know the mechanisms involved, nor the interactive effects of multiple factors, such as warming, UV radiation, nutrient limitation, deoxygenation and cell surface pH (Flynn *et al.* 2012; Milligan and Morel 2002) upon silicate mineralisation.
- (2) Higher  $CO_2$  concentrations appear to favour growth enhancement of larger rather than smaller diatoms (Feng *et al.* 2010) (Y Wu, AJ Irwin, D Suggett, D Campbell, ZV Finkel, unpubl. data). Mechanistic studies are needed to examine responses of differently sized diatoms to ocean acidification, to discriminate among direct size effects (Barton *et al.* 2013) and effects of taxonomic distinctions in cell structures (Mitchell *et al.* 2013) or metabolisms.
- (3) Coastal and pelagic water diatoms may react differently to ocean acidification due to their pre-adaptations to different regimes of mixing, nutrient and diel pH changes. In coastal waters, photosynthetic carbon fixation, and night-time respiration per volume of seawater is much higher, leading to high pH during the day and low pH during the night. Little is known about diatom responses to diel pH changes under elevated CO<sub>2</sub> concentrations as well as to diel changes in the respiration index of the water.
- (4) To guide the scope of studies responses of diatoms to ocean acidification should be examined under expected combinations of environmental changes. Although UV radiation appears not to influence the growth of some diatoms under elevated CO<sub>2</sub> level (Wu et al. 2012a), UV-B (280-315 nm) seems to counteract some effects of high CO<sub>2</sub> and low pH (Li et al. 2012a). Further studies are needed to explore physiological responses of diatoms to reasonable exposures to UV, temperature rise (Shatwell et al. 2012), fluctuation of irradiance to model changing mixing conditions, nutrient limitation and deoxygenation under elevated CO2 and acidification conditions. Hundreds of genes in P. tricornutum are upregulated after acclimation to ocean acidification conditions (Y Li, F Su, Y Wu, KJ Wang, K Gao, unpubl. data), but little has been documented on long-term acclimation of open ocean diatoms, and studies on interactions between ocean acidification and the progression from exponential to stationary phase are just beginning (Orellana et al. 2013).
- (5) Evolutionary responses to increasing CO<sub>2</sub> concentration in diatoms should be examined for hundreds to thousands of generations. The coastal strain *Thalassiosira pseudonana*



**Fig. 2.** Decreased net photosynthetic  $O_2$  evolution rates of algae with increased external partial pressure ratio of  $O_2$  to  $CO_2$  or respiration index (log<sub>10</sub>(pO<sub>2</sub>/pCO<sub>2</sub>): (*a*) diatoms *Thalassiosira pseudonana* and (*b*) *Phaeodactylum tricornutum* and (*c*) green alga, *Ulva prolifera*, grown under elevated (HC, 1000 µatm) or ambient (LC, 390 µatm) CO<sub>2</sub> levels (re-constructed from Gao *et al.* 2012*b*; Xu and Gao 2012).

CCMP 1335 showed little evidence of evolutionary adaptation over months of growth at elevated CO<sub>2</sub> (Crawfurd *et al.* 2011). Evolutionary responses of the freshwater green alga *Chlamydomonas* sp. at high CO<sub>2</sub> demonstrated that some adapted cell lines lost CCM capabilities (Collins and Bell 2004; Collins *et al.* 2006). Lohbeck *et al.* (2012) found that 500 generations of selection at high CO<sub>2</sub> led to recovery of a coccolithophore's growth rates and calcification, although 680 generations of selection at high  $CO_2$  did not show such a trend in *Gephyrocapsa* oceanica (Jin *et al.* 2013*b*).

(6) Monitoring community abundance of diatoms together with other key taxa over longer time scales is important to gain *in situ* information on their responses to environmental changes (Mutshinda *et al.* 2013). These field data obtained from different waters, when combined with mechanistic from controlled experiments would provide valuable insight into future climate change impacts upon phytoplankton communities.

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