

# Effect of nitrogen fertiliser management on soil mineral nitrogen, nitrous oxide losses, yield and nitrogen uptake of wheat growing in waterlogging-prone soils of south-eastern Australia

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**Abstract.** Some of the highest nitrous oxide (N<sub>2</sub>O) emissions arising from Australian agriculture have been recorded in the high-rainfall zone (>650 mm) of south-western Victoria. Understanding the association between nitrogen (N) management, crop N uptake and gaseous losses is needed to reduce N<sub>2</sub>O losses. Field experiments studied the effect of N-fertiliser management on N<sub>2</sub>O emissions, crop N uptake and crop productivity at Hamilton and Tarrington in south-western Victoria. Management included five rates of urea-N fertiliser (0, 25, 50, 100 and 200 kg N/ha) topdressed at either mid-tillering or first-node growth stages of wheat development; urea-N deep-banded 10 cm below the seed at sowing; and urea coated with the nitrification inhibitor DMPP (3,4-dimethylpyrazole phosphate) was either topdressed or deep-banded. Pre-sowing soil profile chemical properties were determined before static chambers were installed to measure N<sub>2</sub>O losses, accompanied by wheat dry matter, crop N uptake and grain yield and quality, to measure treatment differences. N<sub>2</sub>O losses increased significantly ( $P \leq 0.10$ ) where urea-N was deep-banded, resulting in a 2–2.5-fold increase in losses, compared with the nil N control. The high N<sub>2</sub>O losses from deep-banding N appeared to result from winter waterlogging triggering gaseous or drainage losses before wheat reached peak growth and demand for N in spring. Despite the high losses from deep-banding urea-N, grain yields were largely unaffected by N management, except at Hamilton in 2012, where topdressed wheat growing in a soil with large reserves of NO<sub>3</sub><sup>−</sup>-N, and later experiencing post-anthesis water deficit resulted in a negative grain yield response. All sites had high concentrations of soil organic carbon (>2.8%) and the potential for large amounts of N mineralisation throughout the growing season to supplement low N fertiliser recovery. However, topdressed urea-N resulted in significant enrichment of crop tissue ( $P \leq 0.004$ ) and associated positive response in grain protein compared with the deep banded and nil N treatments. 3,4-Dimethylpyrazole phosphate (DMPP)-coated urea provided no additional benefit to crop yield over conventional urea N. Our study highlighted the importance of synchronising N supply with peak crop N demand to encourage greater synthetic N uptake and mitigation of N<sub>2</sub>O losses.

**Additional keywords:** crop nitrogen recovery, 3,4-dimethylpyrazole phosphate, nitrification inhibitor, raised bed, static chamber, water-filled pore space, *Triticum aestivum*.

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## Introduction

Nitrous oxide (N<sub>2</sub>O) has a warming potential ~300 times greater than that of carbon dioxide (CO<sub>2</sub>) (Crutzen 1981) and is considered an important greenhouse gas (IPCC 2013). Atmospheric N<sub>2</sub>O concentrations have risen by ~50 ppb since the beginning of the industrial revolution, and a significant proportion of this increase has been attributed to the loss of N<sub>2</sub>O from agricultural soils (IPCC 2013), largely due to the increased use of synthetic nitrogen (N) fertilisers (Van Groenigen *et al.* 2010). Second to water, N is an essential requirement for crop growth, and as human populations

continue to rise, so too global demand for N fertiliser, forecast to increase by 1.4% annually over the previous year from 2014 to 2018 (FAO 2015). Consequently, global emissions of N<sub>2</sub>O are also likely to increase.

Over the past 20 years, demand for N fertiliser in Australian cropping systems has doubled (Lake 2012), with a shift away from traditional pasture legumes grown in sequence with annual crops to more intensified, continuous cropping with greater reliance on synthetic N (Rovira 1994; Hooper *et al.* 2003; Edwards *et al.* 2012). Unfortunately, N supplied through synthetic sources is likely to produce more N<sub>2</sub>O than N supplied

through the mineralisation of legume residues (Jensen *et al.* 2012; Schwenke *et al.* 2015). Both Barton *et al.* (2008) and Officer *et al.* (2015) showed small but slightly elevated  $\text{N}_2\text{O}$  emissions from the addition of N fertiliser to dryland wheat in semi-arid climatic zones (<450 mm) of southern Australia. In the high-rainfall zone (>650 mm) of south-western Victoria Harris *et al.* (2013) measured large losses of  $\text{N}_2\text{O}$  (<600 g  $\text{N}_2\text{O}$ -N/ha.day) where high inputs of N fertiliser were applied to soils prone to prolonged periods of saturation, with high background levels of mineral N (>200 kg/ha to 100 cm depth) and organic carbon (C, >3.5%).

Although many studies have shown higher rates of N fertiliser produce greater soil-based  $\text{N}_2\text{O}$  losses (Grant *et al.* 2006; Mosier *et al.* 2006; Halvorson *et al.* 2008; Ma *et al.* 2010; Hoben *et al.* 2011; Liu *et al.* 2012), few have attempted to link agronomic N efficiencies with  $\text{N}_2\text{O}$  production (Van Groenigen *et al.* 2010). Best management practices that balance N supply to meet crop demand have been shown to produce lower  $\text{N}_2\text{O}$  emissions (Snyder *et al.* 2007; Scheer *et al.* 2008; Scheer *et al.* 2013; Macdonald *et al.* 2015). Because N fertiliser input is known to influence both crop yield and  $\text{N}_2\text{O}$  loss, it seems reasonable that improving crop N uptake may help to minimise greenhouse-gas emissions from cropping systems. Mosier *et al.* (2006) argued that by linking crop productivity with greenhouse-gas output, it was possible to determine appropriate levels of N fertiliser input for optimising economic return and environmental sustainability.

Another possible strategy for improving crop N utilisation is the use of nitrification inhibitors such as DMPP (3,4-dimethylpyrazole phosphate; BASF Germany), the active ingredient in ENTEC fertilisers. Nitrification inhibitors are designed to delay the oxidation of ammonium ( $\text{NH}_4^+$ -N) to nitrite ( $\text{NO}_2^-$ -N) (Zerulla *et al.* 2001), to keep N in a form less able to escape to the atmosphere as  $\text{N}_2\text{O}$  (Weiske *et al.* 2001; Pfab *et al.* 2012), or leached deeper into the soil profile (Di and Cameron 2012). Greater retention of soil inorganic N implies potentially greater opportunities for plant N uptake (Liu *et al.* 2013). Studies have shown a reduction in  $\text{N}_2\text{O}$  production in response to the application of DMPP (Weiske *et al.* 2001; Belastegui-Macadam *et al.* 2003; Merino *et al.* 2005; Chen *et al.* 2010; Pfab *et al.* 2012; Di and Cameron 2012); however, few have demonstrated improved crop N uptake and grain yield (Liu *et al.* 2013).

This paper examines the role of N-fertiliser management in crop N utilisation, crop productivity and soil-based  $\text{N}_2\text{O}$  losses from wheat growing in waterlogging-prone soils with high concentrations of organic C in south-western Victoria. The fertiliser strategies included different rate, placement (deep-banded v. topdressing) and DMPP-coated N fertiliser, imposed in three separate field experiments; to test the hypothesis that improving crop N fertiliser uptake can reduce  $\text{N}_2\text{O}$  losses from crops growing in waterlogging-prone soil.

## Materials and methods

### Experimental sites

Field experiments were conducted at Hamilton (142°4'15"E, 37°49'27"S) in 2012 and 2014 and at Tarrington (142°6'9"E, 37°47'6"S) in 2013, in south-western Victoria on Ferric-

Eutrophic Brown Chromosol soil (Isbell 2002). Soils at the Hamilton and Tarrington sites were characterised by moderate levels of clay (10–20%) in the topsoil (0–30 cm), but thereafter clay content abruptly increased and remained high (60–70%) throughout the rest of the soil profile.

Before the study, raised beds were formed at the Hamilton site. The site was sprayed with a tank mix of glyphosate (1080 g a.i./ha) and carfentrazone-ethyl (18 g a.i./ha) on 1 September, 5 October and 14 December 2011, to impose a chemical fallow, followed by deep-ripping on 13 January and 7 March and power-harrowing on 15 March. The beds were formed on 21 March 2012. Beds were 1.35 m wide and adjacent furrows 15 cm deep by 35 cm wide. Canola (*Brassica napus* L.) cv. Stingray was sown at 3 kg/ha with a basal application of mono-ammonium phosphate fertiliser (10 kg N and 22 kg phosphorus (P)/ha) with the seed in 2013, the intervening year between experiments. By contrast, the Tarrington site was conventionally flat with a slight slope (1%); canola stubble was incorporated by tillage on 12 March 2013. Winter wheat (*Triticum aestivum* L.) and canola had been rotated at both sites before the study.

### Experimental design

All of the three field experiments comprised a completely randomised block design with seven treatments replicated five times. Plots were 10 m long by 3.4 m wide; or two raised beds with one dividing furrow per plot at the Hamilton sites. The treatments included: a nil N experimental control (0N); four rates of granular urea-N fertiliser topdressed at 10 (TD10N), 35 (TD35N) and 85 (TD85N) and 185 (TD185N) kg N/ha at first-node (Zadoks growth stage Z31) or mid-till (Z25) stage of wheat growth (Zadoks *et al.* 1974); 85 kg N/ha of urea coated with DMPP (ENTEC, Incitec Pivot Ltd, Melbourne), either topdressed at first-node (DMPP85N@Z31) or deep-banded at sowing (DMPP85N@Z00); and 85 kg N/ha of urea-N deep-banded at sowing (DB85N@Z00).

Deep banding involved drilling either urea or DMPP-coated urea to a depth of 10 cm below the intended wheat seeding depth (~5 cm) on the day before planting. All treatments received a basal application of 15 kg P/ha at sowing. With the exception of the 0N treatment, a basal application of 15 kg N/ha was also applied with the seed when wheat cv. Bolac was sown on 31 May 2012 and 7 May 2014 at the Hamilton sites, and 9 May 2013 at the Tarrington site. Fertiliser was topdressed to wheat on 10 September 2012 and 19 August 2013, coinciding with the appearance of the first node at the respective Hamilton and Tarrington sites, and on 18 July 2014 around mid-tillering at Hamilton. Urea and DMPP-coated urea were topdressed by hand to the treatments allocated for topdressing at all sites, and fertiliser was not applied to the adjacent furrows at the Hamilton sites.

### Crop management

Before sowing wheat at the Hamilton and Tarrington sites, a tank mix of glyphosate (1080 g a.i./ha) and carfentrazone-ethyl (18 g a.i./ha) was sprayed to eradicate weeds. Pyroxasulfone (100 g a.i./ha) was sprayed shortly before sowing. Plots were sown by using a cone seeder with knife-points and press-wheels spaced 15 cm apart; furrows were not planted on the raised bed

sites. A basal application of N and P was applied with the seed as di-ammonium phosphate (15 N and 15 kg P/ha) to all treatments except for 0N, which received double superphosphate (15 kg P/ha). Fertiliser was treated with flutriafol (125 g a.i./ha) to combat threats of stripe rust (caused by *Puccinia striiformis*). Control of germinating wild radish (*Raphanus raphanistrum* L.) and capeweed (*Arctotheca calendula*) populations involved the use of MCPA and diflufenican (188 g and 19 g a.i./ha, respectively). Annual ryegrass (*Lolium rigidum*) populations were controlled with tralkoxydim (200 g a.i./ha). Propiconazole (125 g a.i./ha) fungicide was applied to control outbreaks of Septoria leaf blotch (*Mycosphaerella graminicola*). Stubbles remained standing on the sites, over the fallow during the experimental period.

#### *Climatic and soil temperature measurements*

An automated tipping-bucket rain gauge (Hastings Dataloggers, Port Macquarie, NSW) installed in close proximity to the experimental sites measured hourly rainfall at the Hamilton and Tarrington sites. Hourly topsoil water was monitored by theta probes (Theta-Probe MK2x; Delta-T Devices Ltd, Burwell, UK) installed to 6 cm depth. Theta probes were installed in the TD10N treatments at the Hamilton and Tarrington sites, and were placed on top of the raised bed and in the middle of the adjacent furrow at the Hamilton sites.

#### *Soil sample collection for chemical analysis and bulk density*

Five deep soil cores (internal diameter 42 mm) were randomly collected from each replicate of the Hamilton and Tarrington sites before experimentation. On each occasion, the cores were divided into 10-cm increments to 40 cm depth, and thereafter into 20-cm increments. Four of the five cores collected were combined for each layer within each replicate. Samples were then oven-dried at 40°C for 48 h and passed through a 2-mm sieve before chemical analysis. The remaining core collected from each replicate was weighed, oven-dried at 105°C for 48 h and weighed again to determine gravimetric water, bulk density and volumetric water.

Surface soil mineral N ( $\text{NH}_4^+$  and  $\text{NO}_3^-$ ) was measured monthly at the Hamilton and Tarrington sites, with more frequent samples taken before and after N fertiliser applications. On each occasion, 12–15 soil cores (internal diameter 20 mm) were randomly collected to a depth of 10 cm from each plot; at the Hamilton sites, separate soil samples were taken from the beds and from the adjacent furrows. During the Hamilton 2014 experiment, additional soil samples were taken from the 10–20 cm soil layer, from the bed tops over the winter period. Samples were oven-dried at 40°C for 48 h, and passed through a 2 mm sieve in preparation for soil mineral N analysis.

Additional topsoil samples were taken by hand to measure bulk density by using stainless-steel rings (internal diameter 50 mm) to 10 cm depth from each replicate at the Hamilton and Tarrington sites. In each replicate 10 random samples were collected; at the Hamilton sites, separate soil samples were taken from the beds and from the adjacent furrows. Samples were weighed and then oven-dried at 105°C for 48 h before reweighing.

#### *Baseline soil chemical and physical properties*

Analysis of soil before experimentation showed that organic C, total N and P and iron (Fe) concentrations generally declined with depth and that soil pH, exchangeable sodium and bulk density increased with depth at both the Hamilton and Tarrington sites (Table 1). Exchangeable aluminium (Al), soil P and sulphur (S) concentrations were unlikely to inhibit crop growth; however, concentrations of potassium (K) were generally marginal in the surface soil layers at both sites. In the topsoil layer (0–10 cm) of the furrow at the Hamilton site, all soil parameters were lower than in the adjacent beds, with the exception of exchangeable Al and bulk density (Table 1).

#### *Crop measurements*

Crop emergence was measured at the second-leaf stage (Z12) of wheat growth, by counting plants on both sides of a 0.5-m stick, randomly placed 10 times within each plot. Wheat biomass was measured at first node (Z31), anthesis (Z65) and maturity (Z93) by cutting a 1-m row of crop at two random locations within each plot. On each occasion the crop was cut at ground level and bulked within each plot, and subsamples retained and oven-dried at 65°C until constant weight was reached. After drying, spike density was measured by counting fully emerged wheat ears collected from anthesis (Z65) biomass samples, and maturity (Z93) biomass samples were threshed to separate grain from straw to determine respective N concentrations.

Grain yield was measured by mechanically harvesting each plot. A subsample of grain was retained to assess grain quality. Wheat grain protein was calculated by multiplying grain N concentration by 5.7 (Terman 1979). Grain weight was quantified by the mass of 1000 grains, and grain screenings by measuring the proportion of a hectolitre (hL) of grain passing through a 2-mm sieve.

#### *Collection of $\text{N}_2\text{O}$ gas samples*

Concentrations of  $\text{N}_2\text{O}$  gas were measured on a 1–2-week frequency over the growing season, with more intensive samplings around N fertilisation at sowing (Z00) and topdressing (Z25 and Z31). Over the fallow period as topsoil moisture levels declined, gas collection frequency was extended to monthly. Gas was captured in vented static chambers constructed from 25-L PVC drums (internal diameter of 300 mm), with the bases cut off and fitted with one-way valves, rubber septa and battery-powered computer fans to provide continuous air circulation (Harris *et al.* 2013). At sampling, the chambers were placed into *in-situ* ground-level metal troughs filled with water, with each plot containing two troughs; except at the Hamilton sites, where two troughs were installed on the beds and two more in the adjacent furrows. The close proximity of troughs was necessary to place raised platforms between troughs to minimise soil compaction and plot damage from frequent samplings. On all occasions, fluxes were measured between 0900 and 1500 hours. Samples were collected by syringe at 0, 20, 40 and 60 min after chamber emplacement. Air samples (20 mL) were injected into 12-mL evacuated Exetainers (Labco Ltd, High Wycombe, UK) and analysed by gas chromatography. Tinytag Plus 2 temperature data loggers (Hastings Dataloggers, Port Macquarie, NSW)

Table 1. Soil chemical and physical properties at Hamilton and Tarrington sites in south-western Victoria

Soil collected on 21 May 2012 and 28 April 2014 at Hamilton, and 16 April 2013 at Tarrington. Values are mean ( $\pm$  s.e.) of five replicates for each experiment. Separate 0–10 cm samples were collected from the bed top and adjacent furrow at Hamilton

Soil depth (cm)	Organic carbon <sup>A</sup>	Total nitrogen <sup>B</sup> (%)	Soil pH <sup>C</sup> (CaCl <sub>2</sub> )	Exchangeable aluminium <sup>D</sup> (%)	Exchangeable sodium <sup>D</sup> (%)	Sulfur <sup>E</sup>	Colwell phosphorus <sup>F</sup> (mg/kg)	Colwell potassium <sup>F</sup> (mg/kg)	DTPA iron	Bulk density <sup>G</sup> (g/cm <sup>3</sup> )
<i>Hamilton 2012 (bed top)</i>										
0–10	3.10 ( $\pm 0.04$ )	0.32 ( $\pm 0.00$ )	5.24 ( $\pm 0.11$ )	0.32 ( $\pm 0.11$ )	2.44 ( $\pm 0.23$ )	37.54 ( $\pm 1.22$ )	55.40 ( $\pm 3.20$ )	137.80 ( $\pm 10.71$ )	296.41 ( $\pm 16.71$ )	1.06
10–20	2.61 ( $\pm 0.02$ )	0.26 ( $\pm 0.01$ )	5.10 ( $\pm 0.08$ )	0.49 ( $\pm 0.15$ )	2.32 ( $\pm 0.12$ )	24.78 ( $\pm 1.36$ )	40.20 ( $\pm 2.01$ )	105.60 ( $\pm 5.71$ )	287.12 ( $\pm 7.33$ )	1.25
20–30	1.32 ( $\pm 0.11$ )	0.13 ( $\pm 0.00$ )	5.02 ( $\pm 0.06$ )	1.04 ( $\pm 0.27$ )	2.95 ( $\pm 0.14$ )	10.50 ( $\pm 1.59$ )	12.80 ( $\pm 0.73$ )	65.80 ( $\pm 5.43$ )	147.78 ( $\pm 7.65$ )	1.37
30–40	1.02 ( $\pm 0.07$ )	0.10 ( $\pm 0.01$ )	5.46 ( $\pm 0.07$ )	0.33 ( $\pm 0.13$ )	3.66 ( $\pm 0.21$ )	9.74 ( $\pm 1.25$ )	7.60 ( $\pm 2.40$ )	67.40 ( $\pm 11.49$ )	72.17 ( $\pm 14.14$ )	1.39
40–60	0.63 ( $\pm 0.09$ )	0.09 ( $\pm 0.00$ )	5.84 ( $\pm 0.06$ )	0.01 ( $\pm 0.00$ )	5.53 ( $\pm 0.32$ )	18.98 ( $\pm 3.83$ )	3.00 ( $\pm 0.32$ )	68.40 ( $\pm 6.00$ )	19.23 ( $\pm 2.76$ )	1.55
60–80	0.41 ( $\pm 0.03$ )	0.06 ( $\pm 0.00$ )	6.04 ( $\pm 0.02$ )	0.01 ( $\pm 0.00$ )	7.92 ( $\pm 0.52$ )	35.02 ( $\pm 3.58$ )	2.20 ( $\pm 0.20$ )	69.00 ( $\pm 4.34$ )	6.58 ( $\pm 0.96$ )	1.68
80–100	0.28 ( $\pm 0.02$ )	0.04 ( $\pm 0.00$ )	6.02 ( $\pm 0.04$ )	0.01 ( $\pm 0.00$ )	10.01 ( $\pm 0.44$ )	42.54 ( $\pm 4.26$ )	2.00 ( $\pm 0.00$ )	68.60 ( $\pm 4.82$ )	5.16 ( $\pm 0.28$ )	1.70
<i>Hamilton 2012 (adjacent furrow)</i>										
0–10	2.44 ( $\pm 0.08$ )	0.25 ( $\pm 0.01$ )	5.00 ( $\pm 0.07$ )	0.92 ( $\pm 0.19$ )	2.02 ( $\pm 0.12$ )	19.30 ( $\pm 0.59$ )	38.40 ( $\pm 2.56$ )	102.60 ( $\pm 4.48$ )	283.31 ( $\pm 13.63$ )	1.30
<i>Tarrington 2013</i>										
0–10	3.90 ( $\pm 0.21$ )	0.32 ( $\pm 0.01$ )	5.20 ( $\pm 0.09$ )	1.00 ( $\pm 0.25$ )	1.62 ( $\pm 0.07$ )	15.48 ( $\pm 0.86$ )	30.20 ( $\pm 3.54$ )	102.60 ( $\pm 17.08$ )	115.85 ( $\pm 13.31$ )	1.20
10–20	3.01 ( $\pm 0.22$ )	0.26 ( $\pm 0.01$ )	5.22 ( $\pm 0.09$ )	0.80 ( $\pm 0.24$ )	1.62 ( $\pm 0.06$ )	16.58 ( $\pm 1.15$ )	24.40 ( $\pm 2.34$ )	95.60 ( $\pm 19.43$ )	126.65 ( $\pm 4.60$ )	1.25
20–30	2.67 ( $\pm 0.07$ )	0.24 ( $\pm 0.01$ )	5.36 ( $\pm 0.05$ )	0.75 ( $\pm 0.19$ )	2.15 ( $\pm 0.19$ )	21.80 ( $\pm 2.45$ )	20.20 ( $\pm 0.73$ )	73.20 ( $\pm 8.90$ )	97.35 ( $\pm 9.21$ )	1.37
30–40	2.07 ( $\pm 0.26$ )	0.17 ( $\pm 0.02$ )	5.54 ( $\pm 0.05$ )	0.99 ( $\pm 0.17$ )	3.15 ( $\pm 0.29$ )	29.88 ( $\pm 3.62$ )	12.20 ( $\pm 0.36$ )	55.80 ( $\pm 5.54$ )	53.87 ( $\pm 14.35$ )	1.39
40–60	0.86 ( $\pm 0.14$ )	0.07 ( $\pm 0.01$ )	5.98 ( $\pm 0.06$ )	1.41 ( $\pm 0.13$ )	6.07 ( $\pm 0.45$ )	65.12 ( $\pm 2.16$ )	5.00 ( $\pm 0.63$ )	42.00 ( $\pm 1.30$ )	8.90 ( $\pm 2.31$ )	1.55
60–80	0.46 ( $\pm 0.05$ )	0.04 ( $\pm 0.00$ )	6.00 ( $\pm 0.05$ )	1.45 ( $\pm 0.12$ )	8.46 ( $\pm 0.84$ )	76.78 ( $\pm 3.80$ )	5.40 ( $\pm 1.44$ )	40.60 ( $\pm 2.46$ )	2.98 ( $\pm 0.64$ )	1.68
80–100	0.43 ( $\pm 0.03$ )	0.04 ( $\pm 0.01$ )	5.94 ( $\pm 0.07$ )	1.60 ( $\pm 0.08$ )	10.45 ( $\pm 1.02$ )	75.46 ( $\pm 4.37$ )	2.60 ( $\pm 0.24$ )	43.80 ( $\pm 5.44$ )	2.79 ( $\pm 0.63$ )	1.70
<i>Hamilton 2014 (bed top)</i>										
0–10	2.82 ( $\pm 0.10$ )	0.28 ( $\pm 0.01$ )	5.68 ( $\pm 0.17$ )	0.21 ( $\pm 0.09$ )	1.66 ( $\pm 0.10$ )	39.10 ( $\pm 2.71$ )	62.00 ( $\pm 4.34$ )	155.40 ( $\pm 10.54$ )	313.58 ( $\pm 25.08$ )	1.20
10–20	2.16 ( $\pm 0.09$ )	0.22 ( $\pm 0.01$ )	5.34 ( $\pm 0.05$ )	0.41 ( $\pm 0.08$ )	1.54 ( $\pm 0.10$ )	13.20 ( $\pm 1.21$ )	36.20 ( $\pm 2.03$ )	72.40 ( $\pm 6.61$ )	330.25 ( $\pm 15.72$ )	1.25
20–30	1.01 ( $\pm 0.10$ )	0.10 ( $\pm 0.00$ )	5.50 ( $\pm 0.07$ )	0.39 ( $\pm 0.07$ )	2.41 ( $\pm 0.08$ )	13.30 ( $\pm 2.23$ )	13.60 ( $\pm 0.60$ )	49.20 ( $\pm 3.14$ )	136.94 ( $\pm 6.31$ )	1.37
30–40	0.59 ( $\pm 0.06$ )	0.05 ( $\pm 0.01$ )	5.96 ( $\pm 0.05$ )	0.80 ( $\pm 0.08$ )	3.93 ( $\pm 0.11$ )	13.20 ( $\pm 1.42$ )	7.40 ( $\pm 1.72$ )	51.00 ( $\pm 3.99$ )	45.57 ( $\pm 4.33$ )	1.39
40–60	0.59 ( $\pm 0.11$ )	0.05 ( $\pm 0.01$ )	6.12 ( $\pm 0.04$ )	1.18 ( $\pm 0.12$ )	5.87 ( $\pm 0.43$ )	21.08 ( $\pm 3.20$ )	2.80 ( $\pm 0.20$ )	65.20 ( $\pm 5.27$ )	13.48 ( $\pm 2.01$ )	1.55
60–80	0.38 ( $\pm 0.03$ )	0.04 ( $\pm 0.00$ )	6.04 ( $\pm 0.02$ )	1.10 ( $\pm 0.13$ )	7.57 ( $\pm 0.58$ )	33.06 ( $\pm 2.55$ )	2.25 ( $\pm 0.22$ )	69.00 ( $\pm 3.71$ )	5.72 ( $\pm 0.45$ )	1.68
80–100	0.28 ( $\pm 0.04$ )	0.03 ( $\pm 0.00$ )	6.10 ( $\pm 0.03$ )	0.93 ( $\pm 0.08$ )	9.03 ( $\pm 0.45$ )	36.68 ( $\pm 3.24$ )	2.33 ( $\pm 0.26$ )	71.80 ( $\pm 4.24$ )	5.16 ( $\pm 0.97$ )	1.70
<i>Hamilton 2014 (adjacent furrow)</i>										
0–10	2.39 ( $\pm 0.13$ )	0.24 ( $\pm 0.01$ )	5.70 ( $\pm 0.08$ )	0.19 ( $\pm 0.04$ )	1.69 ( $\pm 0.06$ )	26.16 ( $\pm 1.90$ )	50.60 ( $\pm 1.50$ )	172.80 ( $\pm 8.91$ )	319.51 ( $\pm 8.99$ )	1.30

<sup>A</sup>Measured by Walkley and Black (1934) method.

<sup>B</sup>Measured by combustion of air dry soils using LECO combustion analyser.

<sup>C</sup>Measured in 0.01 M CaCl<sub>2</sub> solution at a 1:5 soil: extract ratio using a glass electrode.

<sup>D</sup>Measured in 0.1 M NH<sub>4</sub>Cl/0.1 M BaCl<sub>2</sub> at a 1:10 soil: extract ratio for 2 h, before concentrations determined by inductively coupled plasma.

<sup>E</sup>Measured by Colwell (1965) and Rayment and Higginson (1992) methods.

<sup>F</sup>Measured by Blair *et al.* (1991) method.

<sup>G</sup>Specified in the *Materials and methods*.



were placed inside one chamber in each replicate to monitor changes in air temperature during gas sampling.

#### *Calibration of the theta probe and conversion to water-filled pore space (WFPS)*

When topsoil samples were collected for soil mineral N analysis, a subsample was retained, weighed and oven-dried at 105°C for 48 h, then weighed again to determine gravimetric soil water, before multiplying by bulk density to determine volumetric water (VSW). A regression analysis determined an equation, used to convert raw theta probe water-content data to VSW. Calibration equations used at the Hamilton sites included the following:

$$\text{Beds: VSW} = (0.0097 \times \text{water content}) + 0.0115 \quad (R^2 = 0.96)$$

$$\text{Furrows: VSW} = (0.0072 \times \text{water content}) + 0.0296 \quad (R^2 = 0.96)$$

The calibration equation used at the Tarrington site was

$$\text{VSW} = (0.0006 \times \text{water content}) + 0.0473 \quad (R^2 = 0.84)$$

The WFPS was then determined by dividing VSW by total porosity (Linn and Doran 1984).

#### *Chemical analyses and calculations*

Soil mineral N was determined by extracting soils with 2 M KCl at a soil:solution ratio of 1:10 for 1 h at 25°C, and measuring the concentration of  $\text{NO}_3^-$ -N and  $\text{NH}_4^+$ -N in this extract on a Lachat Flow Injection Analyzer (Lachat Instruments, Milwaukee, WI, USA) (Searle 1984). Grain and plant N concentrations were measured using a LECO CNS 2000 analyser apparatus (LECO Corp., St. Joseph, MI, USA). Gas samples were analysed by a fully automated gas chromatograph (Agilent 7890A; Agilent Technologies Inc., Wilmington, DE, USA) equipped with a micro electron capture detector to quantify  $\text{N}_2\text{O}$  ( $\text{N}_2\text{O}_{(\text{g})}$ ) concentration and then convert to gas density ( $\text{N}_{(\text{g})}$ ) by

$$\text{N}_{(\text{g})} = \text{N}_2\text{O}_{(\text{g})} \times (P \times 2M_w)/(R \times T)$$

where  $P$  is atmospheric standard air pressure of 101.31 kPa,  $M_w$  is the molecular weight of N,  $R$  is the universal gas constant (8.314 J/K.mol), and  $T$  is chamber air temperature (K). Gas density was then adjusted for chamber volume. Fluxes were calculated from the linear increase in gas density in the chamber headspace with time (Barton *et al.* 2008). Cumulative flux was estimated by taking the mean between adjacent gas sampling points in time and multiplying by the number of days between samplings. In the TD35N@Z31 and TD85N@Z31 treatments, cumulative flux was assumed identical to the 0N control in the period before topdressing N fertiliser at the Hamilton 2012 and Tarrington sites.

#### *Statistical analyses*

Separate statistical analyses were performed to determine the effect of N rate, placement (deep banding *v.* topdressing) and fertiliser type (urea *v.* DMPP-coated urea) on topsoil  $\text{NO}_3^-$ -N and  $\text{NH}_4^+$ -N supply,  $\text{N}_2\text{O}$  fluxes and crop productivity from each experiment. Data collected from the beds and furrows at the

Hamilton sites were analysed separately. Treatment differences in topsoil  $\text{NO}_3^-$ -N and  $\text{NH}_4^+$ -N, cumulative  $\text{N}_2\text{O}$  flux and crop productivity were tested by using analysis of variance (ANOVA) appropriate for a completely randomised block design. Logarithmic (base 10) transformations were used to normalise the cumulative  $\text{N}_2\text{O}$  flux data before analyses. All analyses were undertaken using GENSTAT 14 Edition (Lawes Agricultural Trust, Harpenden, UK).

## **Results**

### *Rainfall*

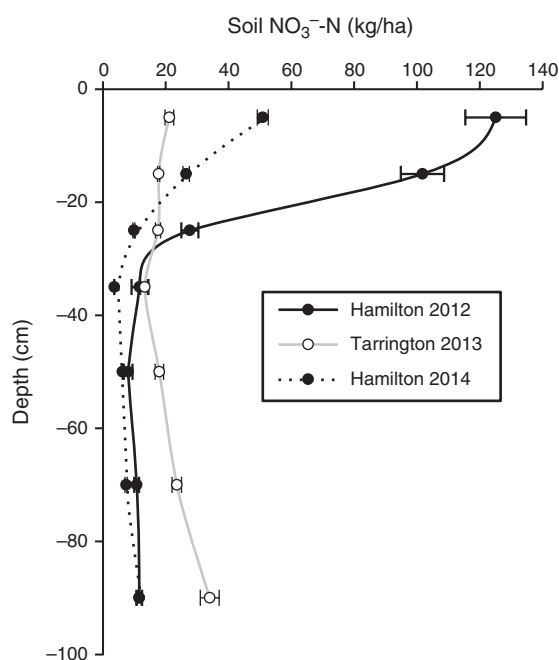
During the 2012, 2013 and 2014 seasons, annual rainfall in the Hamilton region was, respectively, 44, 65 and 187 mm below the long-term (1962–2014) mean of 685 mm. Despite lower-than-average annual rainfall, growing season (April–November) rainfall was within 9 mm of the long-term mean of 538 mm in the 2012 and 2013 seasons, whereas the 2014 season experienced 126 mm less than the long-term mean. Winter rainfall (June–August) was 52 and 22 mm above the long-term average of 233 mm for the 2012 and 2013 seasons, respectively, and 33 mm below average during the 2014 season. Subsequent spring rainfall (September–November) was 50 and 109 mm below the long-term average of 193 mm for the 2012 and 2014 seasons, respectively, and 5 mm above average during the 2013 season.

### *Autumn profile soil $\text{NO}_3^-$ -N at the beginning of each experiment*

Before imposing treatments, the soil profile at the Hamilton 2012, Tarrington 2013 and Hamilton 2014 experimental sites contained (mean  $\pm$  s.e.) 297 ( $\pm 12$ ), 145 ( $\pm 12$ ) and 116 ( $\pm 6$ ) kg soil  $\text{NO}_3^-$ -N/ha, respectively, in the top 100 cm. At the Hamilton 2012 site, 90% of the  $\text{NO}_3^-$ -N was stored in the top 40 cm of the profile, compared with 48% at the Tarrington site and 70% at the Hamilton 2014 site (Fig. 1).

### *Temporal changes in surface soil $\text{NO}_3^-$ -N in each experiment*

At all experimental sites, soil  $\text{NO}_3^-$ -N concentrations increased significantly ( $P < 0.05$ ) after topdressing urea at 85 kg N/ha at either first node (TD85N@Z31) or mid tillering (TD85N@Z25) compared with the 0N control (Table 2). Higher  $\text{NO}_3^-$ -N concentrations persisted under the TD85N treatments throughout the rest of the growing season at all sites and into the fallow at the Hamilton sites in both the 2012–13 and 2014–15 seasons compared with the 0N control. When DMPP was topdressed at first node (DMPP85N@Z31), soil  $\text{NO}_3^-$ -N concentration was significantly ( $P < 0.05$ ) lower in the two samplings after application than in the TD85N@Z31 treatment at the Hamilton 2012 and Tarrington sites (Table 2). The only time that  $\text{NO}_3^-$ -N concentrations were found to be significantly higher under the DB85N@Z00 treatment than the topdressed treatments, or where DMPP was deep-banded at sowing, was when deeper (10–20 cm) samples were collected from the Hamilton site in 2014 (Table 2).



**Fig. 1.** Distribution of soil  $\text{NO}_3^-$ -N (kg/ha) to 100 cm depth in autumn before seeding at the Hamilton 2012, Tarrington 2013 and Hamilton 2014 sites, in south-western Victoria. Bars  $\pm$  s.e. ( $n=5$ ).

#### Temporal changes in surface soil $\text{NH}_4^+$ -N in each experiment

Generally, concentrations of  $\text{NH}_4^+$ -N in the top 10 cm of soil at the first sampling after application were significantly ( $P < 0.05$ ) higher after 85 kg N/ha was topdressed at first node (TD85N@Z31 and DMPP85N@Z31) than in the 0N control at both the Hamilton 2012 and Tarrington sites (Table 3). Thereafter,  $\text{NH}_4^+$ -N concentrations decreased more rapidly under TD85N@Z31 than DMPP85N@Z31 at both the Hamilton 2012 and Tarrington sites (Table 3). In the Hamilton 2014 experiment, significantly ( $P < 0.05$ ) higher concentrations of  $\text{NH}_4^+$ -N were measured in the top 10 cm of soil on 16 June in the deep-banded urea (DB85N@Z00) treatment than in topdressed and 0N treatments, and on 30 July and 15 August in the TD85N@Z25 than all other treatments (Table 3).

#### Temporal changes in WFPS and $\text{N}_2\text{O}$ flux in each experiment

In all experiments, topsoil WFPS responded to periods of high rainfall ( $>15$  mm), associated runoff, surface pondage, evapotranspiration and evaporation. WFPS peaked at 90% on 17 August 2012 in the beds at the Hamilton 2012 site (Fig. 2a), at 108% on 23 August 2013 at the flatter Tarrington site (Fig. 2b), and at 85% on 1 August 2014 in the beds at the Hamilton 2014 site (Fig. 2c).

At all sites, peak  $\text{N}_2\text{O}$  production coincided with the initial stages of the topsoil drying out, after WFPS had peaked in August, with peak flux reaching 372 g  $\text{N}_2\text{O}$ -N/ha.day on 25 September 2012 at the Hamilton 2012 site (Fig. 2a), 65 g  $\text{N}_2\text{O}$ -N/ha.day on 3 September 2013 at the Tarrington site (Fig. 2b),

and 129 g  $\text{N}_2\text{O}$ -N/ha.day on 19 August 2014 at the Hamilton 2014 site (Fig. 2c). At all sites, soil  $\text{N}_2\text{O}$  production was  $<1$  g  $\text{N}_2\text{O}$ -N/ha.day once WFPS dropped below 40% (Fig. 2a, b, c), or when WFPS exceeded 80% (Fig. 2b). At all sites, fluxes were higher where urea-N was either deep-banded (DB85N@Z00) or topdressed (TD85N@Z31 and TD85N@Z25) (Fig. 2a–c) than from the 0N control.

#### Cumulative losses of $\text{N}_2\text{O}$ from each experiment

Cumulative  $\text{N}_2\text{O}$  flux was approximately 2–2.5-fold higher where urea was deep-banded (DB85N@Z00) at sowing than in the 0N control on the beds of the Hamilton sites in 2012 and 2014 (Fig. 3a, c) and the Tarrington site (Fig. 3b). The cumulative  $\text{N}_2\text{O}$  flux from the DB85N@Z00 treatment was significantly ( $P \leq 0.10$ ) greater than from the 0N control at all sites, and from both DMPP deep-banded and topdressed treatments at the Hamilton 2012 and 2014 sites, and some of the topdressed treatments at the Hamilton 2012 and Tarrington sites (Fig. 3). There was no effect of treatment in the adjacent furrows at the Hamilton sites during the 2012–13 and 2014–15 seasons (data not shown).

#### Crop biomass, N uptake, and grain yield and quality

Throughout the three seasons of experimentation, N-fertiliser management had no effect on crop emergence and ear density at anthesis (data not shown), biomass at anthesis and maturity, or harvest index. The exception was biomass measured at first node at the Tarrington site in 2013, where treatments receiving N fertiliser produced 35–47% more biomass than the 0N treatment (Table 4). Grain yield was also unaffected by N-fertiliser management, except at the Hamilton 2012 site, where a significant ( $P = 0.038$ ) incremental decrease in grain yield occurred when urea-N was topdressed compared with the 0N treatment (Table 4). However, at all sites, significant ( $P \leq 0.001$ ) incremental increases in crop N uptake were observed in response to increasing rates of urea-N application (Table 4). The additional crop N uptake resulting from increasing rates of urea-N corresponded to significantly ( $P \leq 0.001$ ) higher grain N concentrations at all sites and significantly ( $P \leq 0.004$ ) higher concentrations of straw N at the Hamilton 2012 and Tarrington sites (Table 4). At all sites, significant ( $P < 0.001$ ) incremental increases in grain protein were observed from increasing urea-N supply (Table 4). However, the positive grain protein responses to increasing urea-N supply corresponded to significantly ( $P \leq 0.04$ ) lower grain weight, and conversely significantly ( $P \leq 0.004$ ) higher grain screenings, in all experiments (Table 4). Furthermore, in all experiments, topdressing DMPP or deep banding N at sowing did not significantly improve wheat yield compared with topdressing urea-N at the equivalent rate (Table 4).

## Discussion

#### Effect of N management on topsoil mineral N supply and associated $\text{N}_2\text{O}$ losses

Placement and/or the timing of urea-N application had a significant effect on  $\text{N}_2\text{O}$  losses; deep-banding urea-N at sowing consistently produced higher losses than the 0N and topdressed urea-N treatments (Fig. 3). Peak crop demand for N

**Table 2. Effect of rate, placement and DMPP-coated N fertiliser on temporal topsoil  $\text{NO}_3^-$ -N concentrations (kg/ha) at Hamilton and Tarrington in south-western Victoria**

0N, Control; DB85N@Z00, DMPP85N@Z00: deep-banded urea, DMPP-coated urea 85 kg N/ha at sowing; DMPP85N@Z31, topdressed DMPP-coated urea 85 kg N/kg at first-node stage; TD35N@Z31, TD85N@Z31: topdressed urea 35, 85 kg N/kg at first-node stage; TD35N@Z25, TD85N@Z25: topdressed urea 35, 85 kg N/kg at mid-tillering stage. n.s., Not significant ( $P > 0.05$ )

Date	0N	DB85N@Z00	DMPP85N@Z31	TD35N@Z31	TD85N@Z31	l.s.d. ( $P = 0.05$ )
<i>Hamilton 2012 0–10 cm</i>						
14 June 12	81	108				n.s.
11 July 12	10	10				n.s.
13 Aug. 12	11	13	11	11	12	n.s.
7 Sep. 12	13	12	14	17	13	n.s.
18 Sep. 12	22	18	21	26	30	8
12 Oct. 12	13	13	19	18	28	7
15 Nov. 12	3	3	4	3	6	2
19 Dec. 12	5	7	8	8	10	3
16 Jan. 13	8	9	17	10	26	9
14 Feb. 13	15	20	26	27	25	8
11 Mar. 13	17	21	29	21	28	6
17 Apr. 13	22	25	37	28	36	6
<i>Tarrington 2013 0–10 cm</i>						
17 May 13	37	37				n.s.
13 June 13	29	36				n.s.
17 July 13	14	19				n.s.
12 Aug. 13	11	15	14	12	13	n.s.
26 Aug. 13	7	11	18	18	33	10
12 Sep. 13	8	10	17	22	33	11
14 Oct. 13	3	4	6	4	6	2
13 Nov. 13	7	9	10	8	10	2
10 Dec. 13	5	7	10	6	11	3
13 Jan. 14	10	11	17	13	15	n.s.
17 Feb. 14	14	17	18	19	20	n.s.
20 Mar. 14	22	27	30	29	28	n.s.
	0N	DB85N@Z00	DMPP85N@Z00	TD35N@Z25	TD85N@Z25	
<i>Hamilton 2014 0–10 cm</i>						
14 May 14	12	14	13	12	13	n.s.
16 June 14	71	116	90	74	88	n.s.
14 July 14	15	24	18	12	14	n.s.
30 July 14	8	6	8	10	23	7
15 Aug. 14	4	6	9	13	48	15
15 Sep. 14	7	6	7	6	27	13
15 Oct. 14	2	3	3	3	18	12
14 Nov. 14	2	3	3	3	16	9
19 Dec. 14	8	10	11	8	20	4
<i>Hamilton 2014 10–20 cm</i>						
16 June 14	57	142	87	72	74	43
14 July 14	23	44	38	27	30	n.s.
15 Aug. 14	5	14	7	6	11	n.s.

in dryland wheat grown in southern Australia occurs during spring, from stem extension until anthesis (Angus 2001). Supplying high rates of urea-N early in the growing season, even when deep-banded in the soil, caused a temporal mismatch between supply and peak crop demand for N, providing greater opportunities for  $\text{N}_2\text{O}$  losses than topdressing in-season around the commencement of stem extension.

Nitrification inhibitors are designed to delay the oxidation of  $\text{NH}_4^+$ -N to nitrite ( $\text{NO}_2^-$ -N) by depressing the activity of soil bacteria (e.g. *Nitrosomonas*) during the nitrification process (Zerulla *et al.* 2001). Addition of DMPP to urea slowed the

conversion of  $\text{NH}_4^+$ -N to  $\text{NO}_3^-$ -N (Tables 2 and 3), consistent with observations of other researchers (Weiske *et al.* 2001; Belastegui-Macadam *et al.* 2003; Merino *et al.* 2005; Chen *et al.* 2010; Di and Cameron 2012; Pfab *et al.* 2012). Slowing the process of nitrification through the use of nitrification inhibitors such as DMPP can reduce subsequent  $\text{N}_2\text{O}$  losses (Weiske *et al.* 2001; Liu *et al.* 2013). The greatest reduction in  $\text{N}_2\text{O}$  emissions observed in our study was when DMPP-coated urea was deep-banded at sowing, compared with deep-banded untreated urea, the former management effectively mitigating emissions during the high-risk winter waterlogging period.

**Table 3.** Effect of rate, placement and DMPP-coated N fertiliser on temporal topsoil  $\text{NH}_4^+$ -N concentrations (kg/ha) at Hamilton and Tarrington in south-western Victoria

0N, Control; DB85N@Z00, DMPP85N@Z00: deep-banded urea, DMPP-coated urea 85 kg N/ha at sowing; DMPP85N@Z31, topdressed DMPP-coated urea 85 kg N/kg at first-node stage; TD35N@Z31, TD85N@Z31: topdressed urea 35, 85 kg N/kg at first-node stage; TD35N@Z25, TD85N@Z25: topdressed urea 35, 85 kg N/kg at mid-tillering stage. n.s., Not significant ( $P>0.05$ )

Date	0N	DB85N@Z00	DMPP85N@Z31	TD35N@Z31	TD85N@Z31	l.s.d. ( $P=0.05$ )
<i>Hamilton 2012 0–10 cm</i>						
14 June 12	5	8				3
11 July 12	15	13				n.s.
13 Aug. 12	13	12	15	15	13	n.s.
7 Sep. 12	10	6	8	8	11	n.s.
18 Sep. 12	24	11	54	23	79	31
12 Oct. 12	10	10	24	12	14	8
15 Nov. 12	4	4	6	7	3	n.s.
19 Dec. 12	7	6	6	6	6	n.s.
16 Jan. 13	7	7	14	7	12	n.s.
14 Feb. 13	7	6	14	6	8	6
11 Mar. 13	13	14	25	12	13	n.s.
17 Apr. 13	5	6	13	6	8	4
<i>Tarrington 2013 0–10 cm</i>						
17 May 13	15	16				n.s.
13 June 13	16	16				n.s.
17 July 13	26	24				n.s.
12 Aug. 13	18	24	16	20	19	n.s.
26 Aug. 13	20	27	72	35	58	20
12 Sep. 13	10	11	25	13	13	6
14 Oct. 13	12	12	19	11	12	4
13 Nov. 13	9	8	11	10	9	n.s.
10 Dec. 13	6	6	6	5	6	n.s.
13 Jan. 14	10	11	10	11	11	n.s.
17 Feb. 14	8	9	9	9	9	n.s.
20 Mar. 14	11	16	11	12	10	n.s.
	0N	DB85N@Z00	DMPP85N@Z00	TD35N@Z25	TD85N@Z25	
<i>Hamilton 2014 0–10 cm</i>						
14 May 14	14	14	17	17	15	n.s.
16 June 14	9	25	21	8	9	13
14 July 14	11	16	18	10	11	n.s.
30 July 14	13	13	13	19	78	34
15 Aug. 14	11	12	10	14	47	15
15 Sep. 14	14	15	20	14	19	n.s.
15 Oct. 14	9	9	10	7	10	n.s.
14 Nov. 14	10	11	11	11	13	n.s.
19 Dec. 14	7	8	7	8	7	n.s.
<i>Hamilton 2014 10–20 cm</i>						
16 June 14	7	21	35	7	9	n.s.
14 July 14	9	12	13	7	8	n.s.
15 Aug. 14	10	10	10	8	10	n.s.

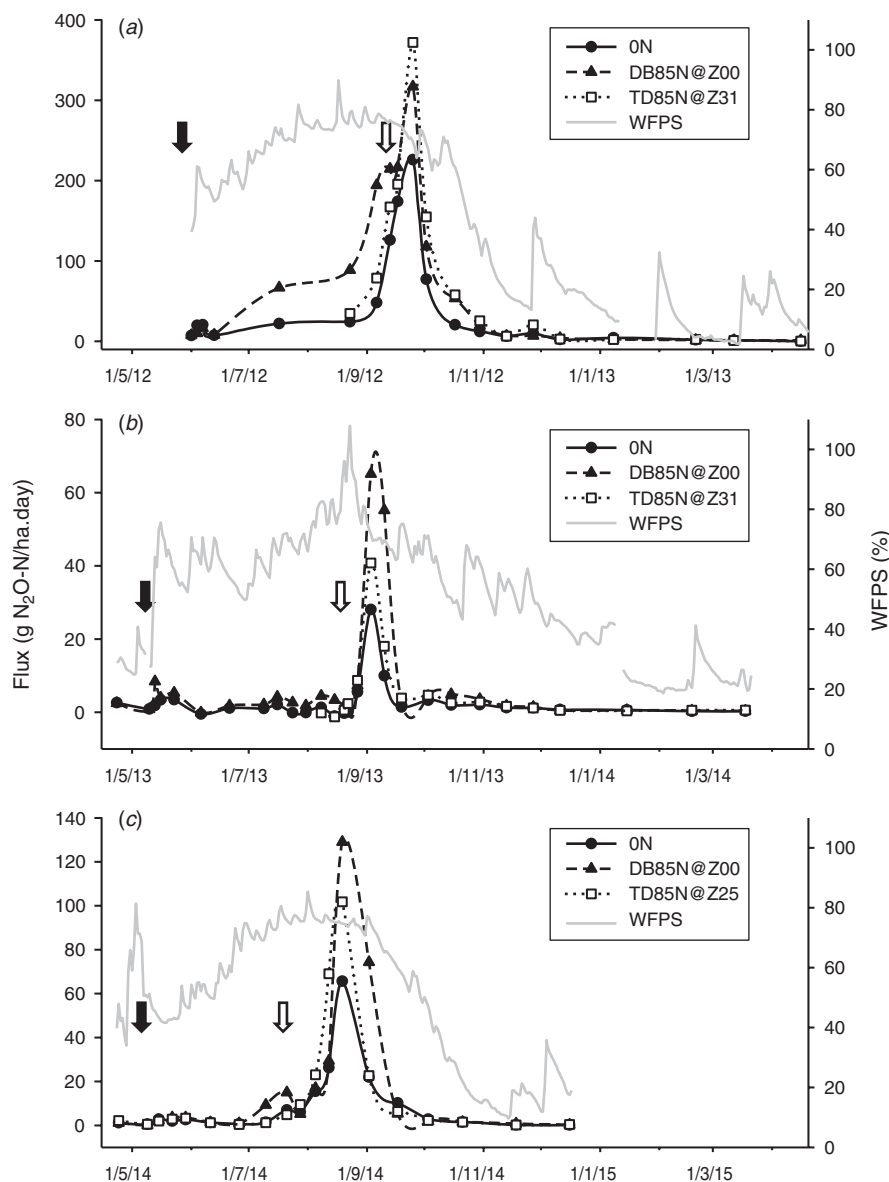
#### *Effect of soil moisture and background soil $\text{NO}_3^-$ on $\text{N}_2\text{O}$ losses*

Soils with abrupt textural changes between the A (topsoil) and B (subsoil) horizons frequently develop surface and subsurface waterlogging caused by rapid saturation of the A horizon from excess winter rainfall and poor infiltration into the B horizon (MacEwan *et al.* 1992). The Brown Chromosol (Isbell 2002) soils used in our studies exhibit strong textural contrast at the junction of the A and B horizons, resulting in prolonged periods of elevated WFPS where raised beds were used at Hamilton in 2012 and 2014 (Fig. 2a, c) and saturated conditions at Tarrington

in the absence of surface drainage (Fig. 2b). Although peak  $\text{N}_2\text{O}$  losses from the soil coincided with the topsoil drying in late winter–early spring (Fig. 2), the  $\text{N}_2\text{O}$  may have formed earlier, before becoming entrapped in the topsoil where it accumulated before release later in the season (Samson *et al.* 1990). Unfortunately, only topsoil (0–10 cm) moisture was monitored in our studies, the placement of theta probes on the junction of the A and B horizons may have enhanced data interpretation.

The magnitude of  $\text{N}_2\text{O}$  losses varied markedly between sites and seasons, due in part to differences in the amounts of soil

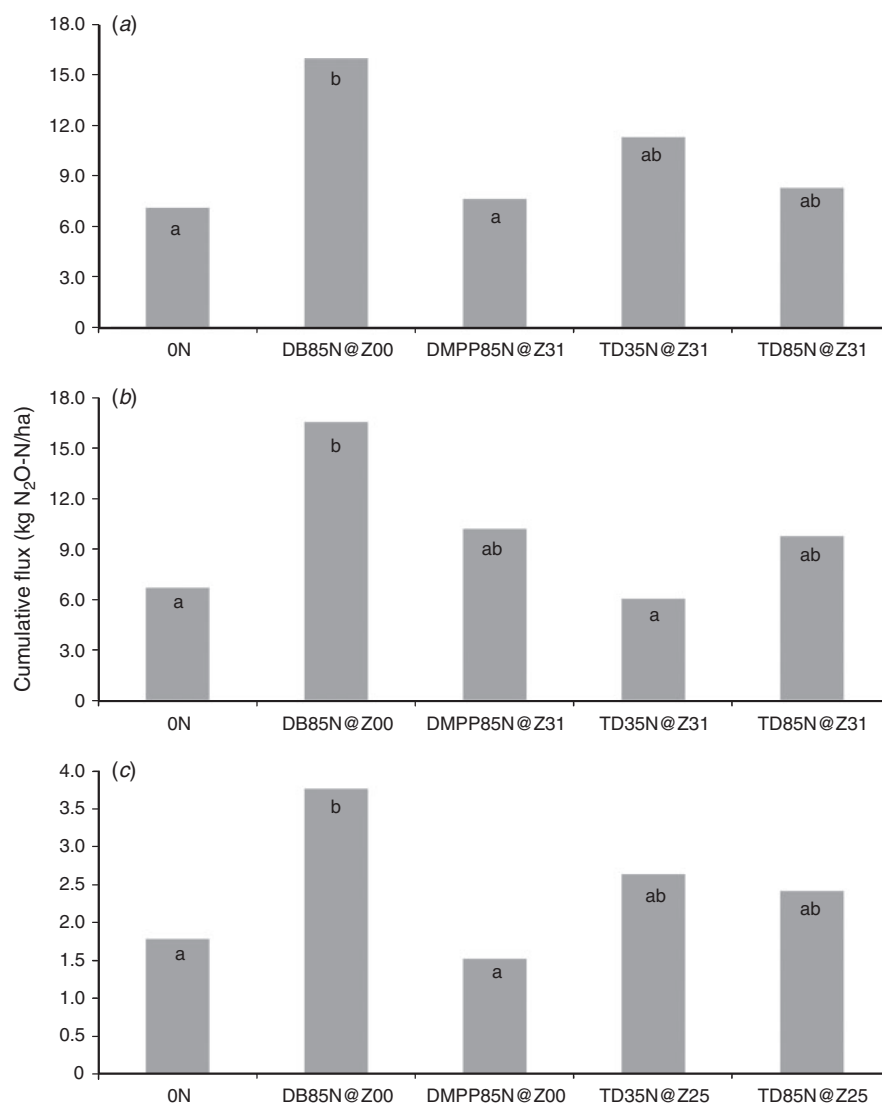




**Fig. 2.** Temporal changes in  $\text{N}_2\text{O}$  flux and water-filled pore space (WFPS) under the control (0N), deep-banded urea 85 kg N/ha at sowing (DB85N@Z00) and topdressed urea 85 kg N/kg at first-node stage (TD85N@Z31) treatments at (a) Hamilton in 2012–13 and (b) Tarrington in 2013–14, and under the 0N, DB85N@Z00 and topdressed urea 85 kg N/kg at mid-tillering stage (TD85N@Z25) treatments at (c) Hamilton in 2014–15 in south-western Victoria. Closed arrow indicates timing of N deep-banded in the DB85N@Z00 treatment, and open arrow indicates N topdressed in the TD85N@Z31 or TD85N@Z25 treatments.

$\text{NO}_3^-$ -N, stored in the soil profile at the beginning of the growing season at the respective sites (Fig. 1). The Hamilton 2012 experiment was imposed after a 6-month chemical fallow, during which soil disturbance occurred when forming raised beds, and in combination with high background soil organic C concentration, this resulted in large amounts of mineral N accumulating before sowing. By contrast, the Tarrington experiment was preceded by ~20 years of annual crops (B. Herrmann pers. comm.), with the previous 2012 canola residue incorporated approximately 2 months before the

experiment was sown; the site also received well below-average summer rainfall. Bending and Lincoln (2000) reported that glucosinolates sourced from residues of brassicas could alter soil microbiota, causing a decline in both  $\text{NH}_3$ -oxidising and nitrifying bacteria, and slowing mineral N transformations for a period. More recently, in laboratory soil incubation experiments, Begum *et al.* (2014) also showed initial N immobilisation and associated reductions in  $\text{N}_2\text{O}$  emissions from the addition of canola (*Brassica napus* L.) residue; however, later in the experiment, the low C : N ratio



**Fig. 3.** Effect of rate, placement and DMPP-coated N fertiliser on cumulative flux over the growing season at (a) Hamilton 2012, (b) Tarrington and (c) Hamilton 2014 sites in south-western Victoria. Treatment means with the same letter are not significantly different after log-transforming the data ( $P < 0.05$  at Hamilton in 2012–13 and 2014–15, and  $P = 0.10$  at Tarrington in 2013–14). N<sub>2</sub>O flux was measured during: (a) 1 June 2012–8 January 2013, (b) 23 April 2013–15 January 2014, (c) 24 April 2014–16 December 2014.

of canola residue produced greater N<sub>2</sub>O losses. Given the absence of a significant yield response to urea-N application at the Tarrington site (Table 4), we speculate that our wheat crop may have accessed a significant proportion of its N supply from mineralisation of the previously incorporated canola residue. Furthermore, frequent wetting and drying of the topsoil (Fig. 2b) over spring at Tarrington in the presence of potentially high amounts of soluble C could have been ideal for rapid mineralisation of fresh residue (Fillery 2001). At the Hamilton site in 2014, the previous surface canola residue was removed before experimentation, possibly reducing immobilisation of NO<sub>3</sub><sup>-</sup>-N.

Differences in the magnitude of N<sub>2</sub>O losses between sites and seasons may also be due to the drainage characteristics of

the respective sites. The N<sub>2</sub>O:N<sub>2</sub> ratios associated with denitrification are strongly influenced by WFPS percentage (Ciarlo *et al.* 2007). At the wetter Tarrington site, WFPS levels exceeded 80% for longer, and therefore, conditions were more conducive to greater N<sub>2</sub> production compared with the raised beds at the Hamilton sites, where WFPS rarely exceeded 80% in both the 2012 and 2014 growing seasons (Fig. 2). However, this theory assumes that denitrification was the source of the N<sub>2</sub>O production, and we acknowledge that dissimilatory NO<sub>3</sub><sup>-</sup>-N reduction, another possible pathway for N<sub>2</sub>O losses, may be occurring in our soils with high concentrations of C (Yin *et al.* 2002), and if so, N<sub>2</sub> gas losses may be less consequential. Although our research has quantified losses of N<sub>2</sub>O under different N-fertiliser

**Table 4. Effect of rate, placement and DMPP coated N fertiliser on aboveground biomass, straw and grain N concentration, N uptake, yield and grain quality at Hamilton and Tarrington in south-west Victoria**

0N, Control; DB85N@Z00, DMPP85N@Z00: deep-banded urea, DMPP-coated urea 85 kg N/ha at sowing; DMPP85N@Z31, topdressed DMPP-coated urea 85 kg N/kg at first-node stage; TD10N@Z31, TD35N@Z31, TD85N@Z31, TD185N@Z31: topdressed urea 10, 35, 85, 185 kg N/kg at first-node stage; TD10N@Z25, TD35N@Z25, TD85N@Z25, TD185N@Z25: topdressed urea 10, 35, 85, 185 kg N/kg at mid-tillering stage. All treatments received 15 kg N/ha applied with the seed at sowing, except the 0N treatment. n.s., Not significant ( $P > 0.05$ ). Values are the means of five replicates

Treatment	N rate	Biomass (kg DM/ha)			Nitrogen (%)		Crop N uptake <sup>A</sup> (kg N/ha)	Grain yield <sup>B</sup> (kg/ha)	Harvest index <sup>C</sup>	Grain protein (%)	1000-grain weight (g)	Grain screenings (%)
		Z31	Z65	Z93	Straw	Grain						
Hamilton 2012												
0N	0	1997	10 967	16 585	0.64	1.93	194	6871	0.36	11.18	33.98	6.75
TD10N@Z31	25	2359	12 227	16 463	0.75	2.12	218	6995	0.37	11.33	31.49	7.37
TD35N@Z31	50	2757	11 807	16 439	0.8	2.21	223	6461	0.35	12.14	29.3	8.93
TD85N@Z31	100	2387	12 547	16 097	0.84	2.47	239	6319	0.35	13.12	30.59	10.05
TD185N@Z31	200	2313	10793	16 107	0.94	2.74	260	6095	0.33	13.84	27.36	11.76
DMPP85N@Z31	100	2401	11 273	16 518	0.8	2.36	233	6463	0.34	12.56	30.69	8.6
DB85N@Z00	100	2611	11 547	15 697	0.8	2.27	226	6964	0.39	11.9	28.65	7.89
F-prob.		n.s.	n.s.	n.s.	<0.001	<0.001	<0.001	0.032	n.s.	<0.001	0.006	<0.001
l.s.d. (P=0.05)					0.094	0.2	21	608		0.59	3.11	1.82
Tarrington 2013												
0N	0	572	6298	10 773	0.41	1.96	139	6100	0.50	11.19	36.71	1.98
TD10N@Z31	25	956	7184	11 109	0.47	1.99	146	6142	0.49	10.9	35.27	1.69
TD35N@Z31	50	890	7186	11 527	0.5	2.02	153	6308	0.48	11.33	36.95	1.82
TD85N@Z31	100	1089	6770	11 528	0.69	2.26	176	6212	0.48	12.14	34.45	2.02
TD185N@Z31	200	1033	7168	11 187	0.8	2.54	194	6150	0.49	13.34	32.57	2.59
DMPP85N@Z31	100	999	6827	10 958	0.62	2.23	167	6170	0.50	12.35	32.8	2.1
DB85N@Z00	100	1073	7154	13 197	0.56	2.15	176	6488	0.43	11.56	35.12	1.66
F-prob.		292	n.s.	n.s.	0.004	0.001	<0.001	n.s.	n.s.	<0.001	0.04	0.004
l.s.d. (P=0.05)					0.18	0.23	20			0.76	2.96	0.42
Hamilton 2014												
0N	0	1356	11 138	14 977	0.28	1.61	127	6432	0.38	9.18	37.22	1.83
TD10N@Z25	25	1631	12 447	15 796	0.37	1.65	143	6618	0.37	9.39	35.59	2.48
TD35N@Z25	50	1887	11 832	15 903	0.29	1.79	145	6523	0.36	10.19	35.64	2.58
TD85N@Z25	100	1598	12 324	15 139	0.31	1.92	153	6577	0.38	10.96	34.89	2.94
TD185N@Z25	200	1594	11 710	14 985	0.44	2.2	173	6135	0.36	12.55	33.95	4.5
DMPP85N@Z00	100	1711	12 691	15 033	0.35	1.84	150	6540	0.38	10.47	35.02	2.41
DB85N@Z00	100	1573	10 972	15 241	0.33	1.82	148	6539	0.38	10.36	33.88	2.91
F-prob.		n.s.	n.s.	n.s.	n.s.	<0.001	0.001	n.s.	n.s.	<0.001	0.038	0.002
l.s.d. (P=0.05)						0.14	17			0.8	2.03	1.08

<sup>A</sup>Grain and straw.

<sup>B</sup>Grain yield at 12% moisture.

<sup>C</sup>Grain yield adjusted to 0% moisture.

management, we remain uncertain about the origin of the N<sub>2</sub>O and, therefore, total gaseous N losses and the associated agronomic implications for high-rainfall cropping soils with high background soil C supply.

#### *Effect of management on crop N-fertiliser recovery and productivity*

Both Edwards (1992) and Anderson *et al.* (1992) showed that establishment of early-sown, well-fertilised, dense, vigorously growing wheat crops on texture-contrast soils was an important strategy for optimising grain yields. Passioura (1992) suggested two main reasons for the success of this strategy. First, it allows roots more time to colonise the subsoil before the onset of winter waterlogging where anoxic conditions were less likely to

impede growth, even though the topsoil might be saturated above. Second, early growth promotes greater utilisation of soil NO<sub>3</sub><sup>-</sup>-N after opening autumn rains stimulate the flush of organic matter mineralisation. In our studies, only a small benefit of N fertiliser, applied with the seed or deep-banded at sowing, was observed at the Tarrington site, and this was evident only in the dry matter data collected at first node (Table 4). Thereafter, no biomass or grain yield response to this added N was found. This finding contradicts the conclusions reached by Watson *et al.* (1976), Drew *et al.* (1979) and Huang *et al.* (1994), who reported that additional N at sowing, or increased supply of N to waterlogged crops, could increase yield and mitigate the harmful effects of waterlogging. Bell *et al.* (2013) reported a critical soil test value at the beginning of the growing season for wheat in Victoria of 62 kg NO<sub>3</sub><sup>-</sup>-N/ha to 60 cm depth; below

this concentration, wheat was likely to respond to additional synthetic N supply. Although the value from Bell *et al.* (2013) is based mainly on drier Victorian environments with much lower yield potentials and associated demand for N, the  $\text{NO}_3^-$ -N concentrations to 60 cm depth in our studies approached this critical concentration only at the Tarrington site (Fig. 1b).

In our experiments, the wheat crops must have had access to more N than just the  $\text{NO}_3^-$ -N stored in the profile around seeding. The large losses of  $\text{N}_2\text{O}$  observed in our studies, by southern Australian cropping standards (Barton *et al.* 2008; Officer *et al.* 2015) and in the absence of any positive grain yield response, seem a contradictory outcome. In-crop N mineralisation has been shown to be an important source of non-synthetic N supply to crops in medium-high-rainfall zones of southern New South Wales (Angus *et al.* 1998). Although soil mineralisation rates were not measured in our studies, applying the model developed by Baldock (2003), accounting for high organic C concentrations at the Tarrington (3.90%) and Hamilton (3.10–2.82%) sites, we estimate that 99–195 kg N/ha may have mineralised over the respective growing seasons. The high in-crop mineralisation potential at both the Hamilton and Tarrington sites partly explains why yields remained unaffected despite the concurrent large losses of  $\text{N}_2\text{O}$  (Fig. 3).

Mineralisation of organic matter only partly explains the outcomes of N-fertiliser management on crop productivity. We estimate the water use efficiency (WUE) at the Hamilton 2012, Tarrington 2013 and Hamilton 2014 sites, respectively, to be 16, 15 and 20 kg grain/mm plant-available water. The benchmark of 20 kg grain/mm plant-available water suggested by French and Schultz (1984) was reached only in 2014. We cautiously rule out other potential nutritional constraints, with S and K concentrations above levels considered to limit yield (Anderson *et al.* 2013; Brennan and Bell 2013), although K appeared marginal in the lower layers of the topsoil (Table 1). Research elsewhere has shown that applications of K to waterlogged wheat can significantly increase yield (Belford *et al.* 1992). Likewise, we cannot rule out that ion (e.g. Fe) toxicities often associated with transient waterlogging (Khabaz-Saberi *et al.* 2014) may have played some role in reducing WUE.

The large reduction in WUE at Tarrington in 2013 appeared largely due the site becoming inundated with water in mid-August (Fig. 2b), reducing soil oxygen diffusion (Bollmann and Conrad 1998) and causing crop damage from transient waterlogging. McDonald and Gardner (1987), MacEwan *et al.* (1992) and Harris *et al.* (2013) have collectively reported reductions in grain yields from 25% to 85% compared with crops grown under improved drainage in south-western Victoria. McDonald and Gardner (1987) observed that waterlogging resulted in shoot N concentrations declining to low levels between stem elongation and anthesis. In our study, topdressing urea-N at around first node resulted in greater crop uptake of N, but had no effect on anthesis and maturity biomass or grain yield (Table 4). Although topdressing N at around first node reduced  $\text{N}_2\text{O}$  emissions, the application to cv. Bolac, a slow-maturing spring-type wheat, may have been too late to influence yield; instead, the extra N was directed into the kernel, thereby increasing grain protein concentrations (Fischer *et al.* 1993).

Surface waterlogging was less of a constraint to crop yield on the raised beds at the Hamilton sites, because the beds are designed to drain excess winter rainfall into the adjacent furrow (Bakker *et al.* 2005). However, yellowing of older leaves was still observed in mid-late winter in both years at Hamilton, symptoms associated with transient N deficiency (Robertson *et al.* 2009), raising the possibility of subsoil waterlogging at the junction of the A and B horizons below the depth of the furrow. The topdressing of urea-N resulted in greater crop N uptake and the alleviation of temporary N deficiency; however, a positive yield response was not observed, despite improved surface drainage. Although our experiments were conducted in three different seasons, anthesis and maturity biomass were greater and conversely harvest index (HI) was much lower on the beds at Hamilton than the undrained Tarrington site. Passioura (1977) illustrated how HI can be a function of water used after anthesis, when biomass is partitioned into grain. Later, Fischer (1979) described a physiological framework for dryland wheat yields in Australia, suggesting 'that grain number in the wheat crop is largely determined by the end of anthesis', but that grain weight appeared largely determined by post-anthesis grain-filling conditions. The low HI at the Hamilton sites suggests post-anthesis water stress resulting from below-average spring rainfall in both seasons, and perhaps excessive pre-anthesis growth from improved surface drainage, contributing to the crops experiencing water deficit late in the growing season.

The higher N concentration in response to the extra urea-N topdressed at around first node to the 2012 crop, along with large reserves of soil  $\text{NO}_3^-$ -N (Fig. 1a), probably resulted in greater accumulation of stem and leaf protein and an associated reduction in water-soluble carbohydrates (van Herwaarden *et al.* 1998a). Although no measurements of kernel density were made in our experiments, van Herwaarden *et al.* (1998b) also showed no significant increase in anthesis and maturity biomass yet a significant increase in spike and kernel density when they applied increasing rates of N fertiliser to wheat growing at Wagga Wagga. The van Herwaarden *et al.* (1998b) crop subsequently experienced water deficit over the grain-filling period and a shortfall of assimilate for translocation into the kernel. The grain yield reduction and associated loss of grain quality, described by van Herwaarden *et al.* (1998b) as 'haying-off', was the likely mechanism for the grain yield reduction observed at the Hamilton 2012 site.

Even less spring rainfall fell at Hamilton in 2014, and yet haying-off was less pronounced with marginal penalty to grain quality, and WUE was much higher than previous seasons. The 2014 season was also characterised by a drier winter, especially during August, with topsoil drying and release of  $\text{N}_2\text{O}$  observed earlier than in 2012 (Fig. 2a, b). Perhaps a longer duration between topsoil drying out and anthesis provided more time for crop root penetration into the subsoil and greater access to subsoil moisture compared with the 2012 crop. Alternatively, the 2014 crop was sown earlier followed by unseasonably warm conditions during May, which may have led to deeper root penetration into the subsoil before the onset of waterlogging. Passioura (1992) suggested that despite roots in the waterlogged soil layers losing their ability to extract water and nutrients, the cortex of the anoxic roots may still act as a conduit for moving shoot-produced assimilates to active roots in the subsoil, and in



turn water and nutrients from the subsoil to the shoot. Kirkegaard *et al.* (2007) highlighted that 'relatively small amounts of subsoil water can be highly valuable to grain yield', reporting efficiencies from subsoil water use three times that typically expected for total seasonal water use. Although we remain uncertain about when the roots colonised the subsoil in 2014, access to deeper subsoil moisture is the only reasonable explanation for the wheat production and grain quality achieved. Perhaps the later sown 2012 crop, with greater N supply and a longer duration of subsoil winter waterlogging confining roots largely to the topsoil, resulted in more severe water deficit at grain fill than that experienced by the 2014 crop.

The absence of a grain yield response to urea-N may not be entirely explained by high rates of organic matter mineralisation, waterlogging, water deficits during grain filling or timing of N application. Riffkin *et al.* (2012) reported significant grain-yield increases from applying high rates of N fertiliser to early-sown, late-maturing winter-type canola at Hamilton during the 2008 growing season. Although canola is a different crop from wheat, longer term results from National Variety Trials testing of wheat varieties in south-western Victoria have also shown longer seasoned winter types out yielding spring-type wheats (<http://www.nvtonline.com.au/google-maps/2014/WheatEaSouth%20West.pdf>). Early-sown, longer season winter-type crops possessing longer vegetative growth phases are likely to utilise more radiation (Riffkin *et al.* 2012), potentially resulting in higher demands for synthetic N and associated yield responsiveness in the south-western Victorian environment.

Yield and grain quality were unaffected by the use of DMPP-treated urea, as there was no evidence of increased retention of the applied N for crop N uptake (Table 4). Furthermore, although nitrification was slowed, no subsequent delayed increase in soil  $\text{NO}_3^-$ -N concentrations was observed after the application of DMPP (Table 2). Soares *et al.* (2012) reported that the nitrification inhibitor dicyandiamide (DCD) coated to urea increased ammonia ( $\text{NH}_3$ ) volatilisation by 5–16%; this was caused by maintaining higher soil  $\text{NH}_4^+$ -N and pH for a longer time than with urea. Although our study involved the use of a different type of nitrification inhibitor, we speculate that higher  $\text{NH}_3$  losses from the use of DMPP may have resulted in no greater net retention of N for crop uptake or subsequent improvements in yield or grain protein over the use of conventional urea (Table 4).

## Conclusion

This study has shown that synchronising N fertiliser supply with peak crop demand can significantly reduce  $\text{N}_2\text{O}$  losses from waterlogging-prone soils. Excessive N supplied early in the growing season provides greater opportunity for N loss over winter, because wheat utilises most of the available N over the early–mid-spring period. Although the supply of N at around the commencement of stem extension enhanced crop N uptake, no improvement in grain yield was found, likely from high rates of in-season organic matter mineralisation, and/or crops encountering water deficit during grain fill, especially during the 2012 growing season. However, the enhanced crop N uptake from urea topdressed around stem extension is likely to benefit

grain quality, especially protein, provided no water deficit is experienced during grain fill. Although DMPP reduced  $\text{N}_2\text{O}$  losses, especially when applied at sowing, there was no evidence of improved yield or grain quality over the use of conventional urea N.

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