
Supplementary material

Abrupt regime shifts in post-fire resilience of Mediterranean mountain pinewoods are fuelled by land use

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Supplementary material S1. Paleoecological methods

The modern vegetation of the peat bog includes *Caricetum carpetanae* communities dominated by *Carex nigra*, *C. echinata* and *Parnassia palustris* (Sánchez-Mata 1989). The vegetation of the area combines Maritime pine forests (*Pinus pinaster*), and *Pinus nigra* subsp. *salzmannii* woodlands with *Genista cinerascens* and *Festuca elegans* subsp. *merinoi* (López-Sáez *et al.* 2016). Overall, the high-supramediterranean belt has the lowest vegetation cover, with a predominance of grasslands, *Juniperus communis* subsp. *alpina* and broom communities (*Cytisus oromediterraneus*, *Echinospartum ibericum*), as well as scattered *Pinus sylvestris* and *P. nigra* subsp. *salzmannii* stands.

A 50-cm sediment core was extracted from “Arroyo de Aguas Frías” peat bog in October 2000 using a Russian corer with a diameter of 50 mm. Peat sections were individually sealed and stored at 4°C prior to laboratory sub-sampling at 1 cm intervals. Five peat samples were ¹⁴C dated using AMS technique (Table S1). The AMS dating was conducted at the Centro Nacional de Aceleradores (CNA) Laboratory (Sevilla, Spain). Radiocarbon dates were calibrated in years cal AD with the CALIB 7.10 software using the calibration curve IntCal13 (Reimer *et al.* 2013). Dates are expressed as intercepts with 2σ ranges. An age-depth model (Fig. S1) was produced using Clam 2.2 software (Blaauw 2010). The best fit was obtained applying a smoothing spline to the available radiocarbon dates. Pollen analysis was carried out on 50 sub-samples of 1 cm³ along the core length. Standard pollen extraction techniques were employed (Faegri and Iversen 1989). Pollen counts of up to 400 grains total land pollen (TLP) per sample were identified and counted. Percentages were calculated based on TLP sum excluding wetland taxa and non-pollen palynomorphs (NPPs). Pollen diagrams based on these data were produced using Tilia and Tilia-Graph v. 2.0.b.5 software (Grimm 1992, 2004). Pollen grains, spores and NPPs were identified using palynological keys and photo atlases (Moore *et al.* 1991; Reille 1999; Cugny *et al.* 2010). Additionally, *Pinus pinaster* was identified according to Carrión *et al.* (2000). Local pollen assemblage zones (LPAZs) were defined on the basis of agglomerative cluster analysis of incremental sum of squares (Coniss) with square root transformed percentage data (Grimm 1987). The number of statistically significant zones was determined using the broken-stick model (Bennett 1996).

Contiguous 1 cm³ sub-samples were retrieved at 1 cm intervals for macroscopic charcoal analysis. Samples were soaked in a 10% NaOH solution for 24 h to disaggregate organic silts, then in a 6% H₂O₂ solution (24 h) to bleach the remaining non-charcoal organic material and thus make charcoal identification easier (Carcaillet *et al.*, 2001). Samples were

gently wet sieved through a 150 μm mesh. The total number of macrocharcoal particles was counted using a stereomicroscope in order to estimate past changes in fire activity at a local scale (Whitlock and Larsen 2001, Higuera *et al.* 2007, Brown *et al.* 2013). Charcoal accumulation rate (CHAR) was calculated by sedimentation rate (cm year^{-1}) and is expressed in particles $\text{cm}^{-2} \text{year}^{-1}$ (Long and Whitlock 2002).

Table S1. Accelerator Mass Spectrometry radiocarbon (AMS- ^{14}C) data with 2σ range of calibration from “Arroyo de Aguas Frías” peat bog.

Laboratory code	Depth (cm)	AMS ^{14}C age BP (yrs.)	Age cal AD	Mean Age cal AD
CNA-111	50	230 ± 30	1532-1950	1737
CNA-112	37	143 ± 20	1669-1944	1802
CNA-113	33	75 ± 30	1691-1923	1847
CNA-114	25	67 ± 20	1696-1919	1868
CNA-115	19	10 ± 30	1690-1960	1900

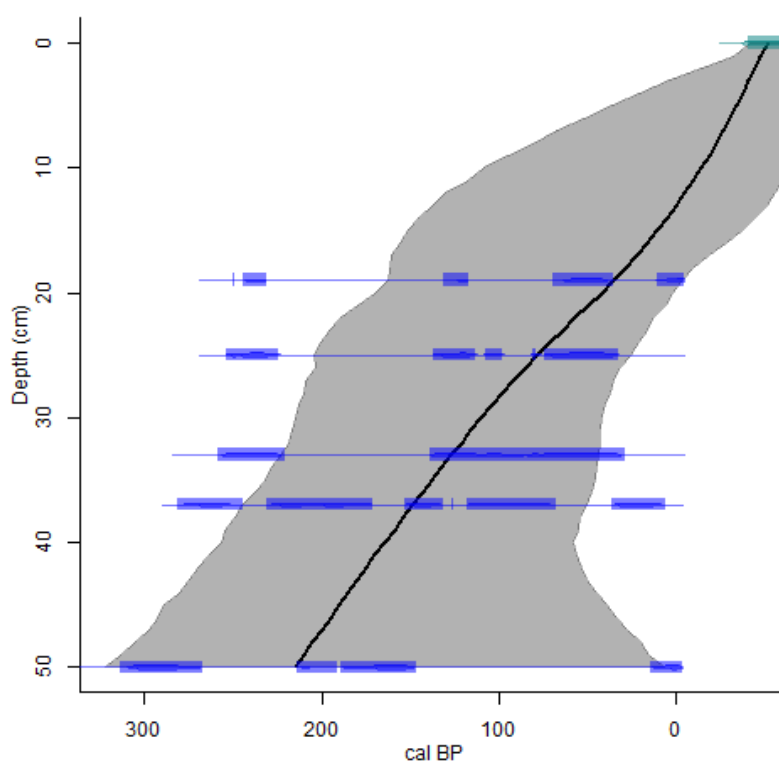


Fig. S1. Age-depth model from “Arroyo de Aguas Frías” peat bog.

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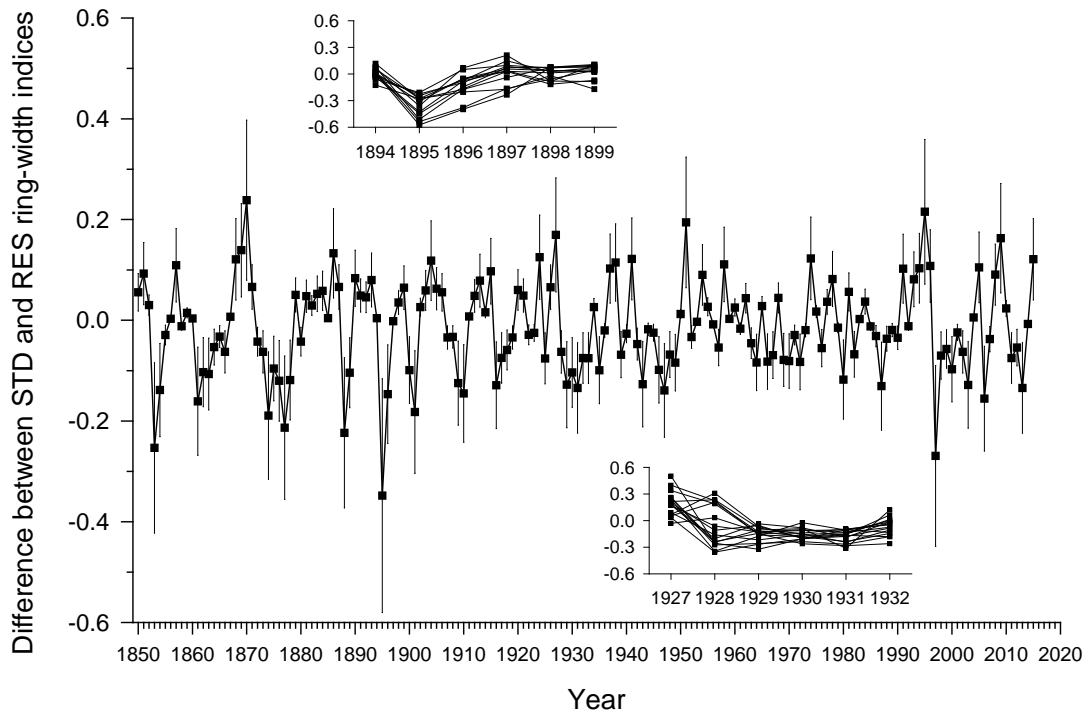


Fig. S2. Differences between the standard and residual ring-width indices observed in site TO (values are means \pm standard errors) and details for two selected periods (insets) when historical fire frequency increased: 1894-1899 and 1927-1932.

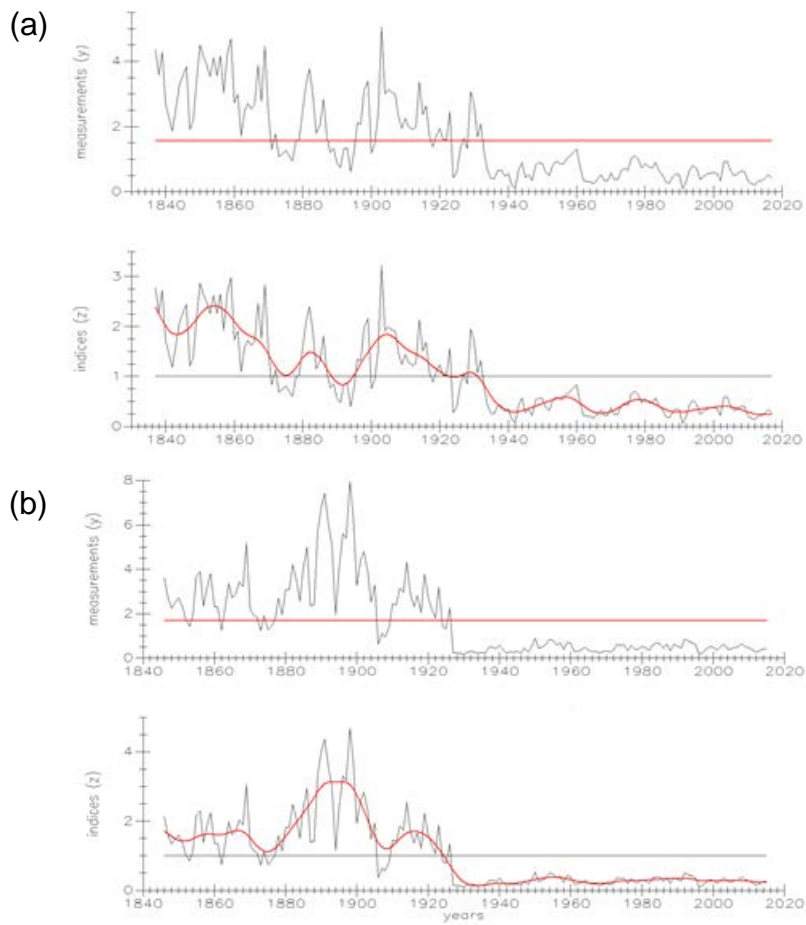


Fig. S3. Examples of trees sampled at sites CE (a) and TO (b) showing punctual and long-term growth declines in the 1890s and late 1920s, respectively. In each plot, the upper series show tree-ring width data (measurements in mm) and a fitted horizontal line, and the lower series show ring-width indices with a polynomial (spline) function.

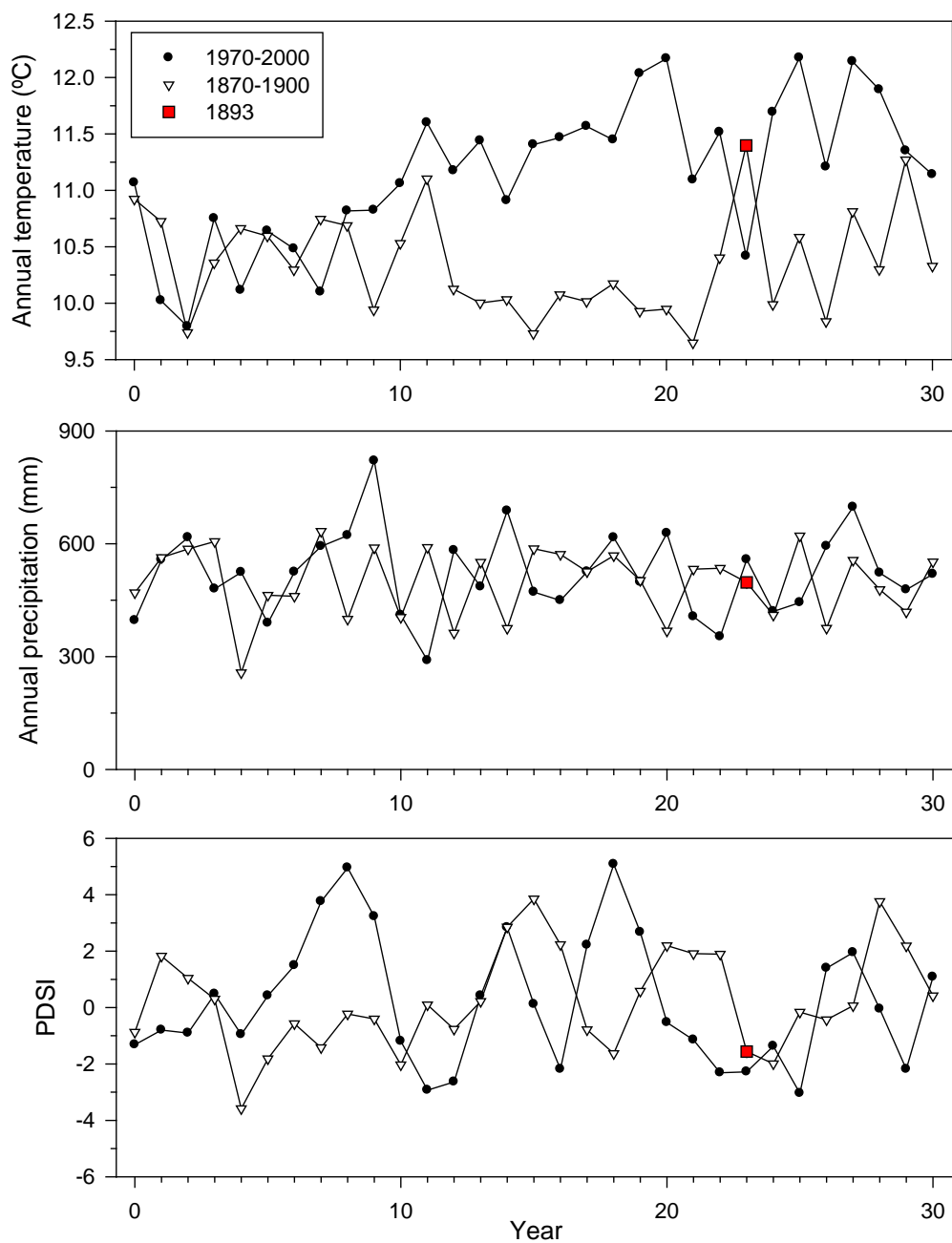


Fig. S4. Reconstructions of annual climate variables for the study area (means or total for the two 0.5° grids with coordinates 4.5° - 5.0° W, 40.0° - 40.5° N and 5.0° - 5.5° W, 40.0 - 40.5° N) considering the last 31 years of the 19th (1870-1900) and 20th centuries (1970-2000). The square corresponds to 1893 when a high fire activity was observed according to historical records and paleoecological proxies (high charcoal accumulation rates, narrow tree rings). Temperature and precipitation reconstructions were produced by Luterbacher et al. (2004) and Pauling et al. (2006), respectively. The Palmer Drought Severity Index (PDSI) was obtained from the Old World Drought Atlas (Cook et al., 2015).

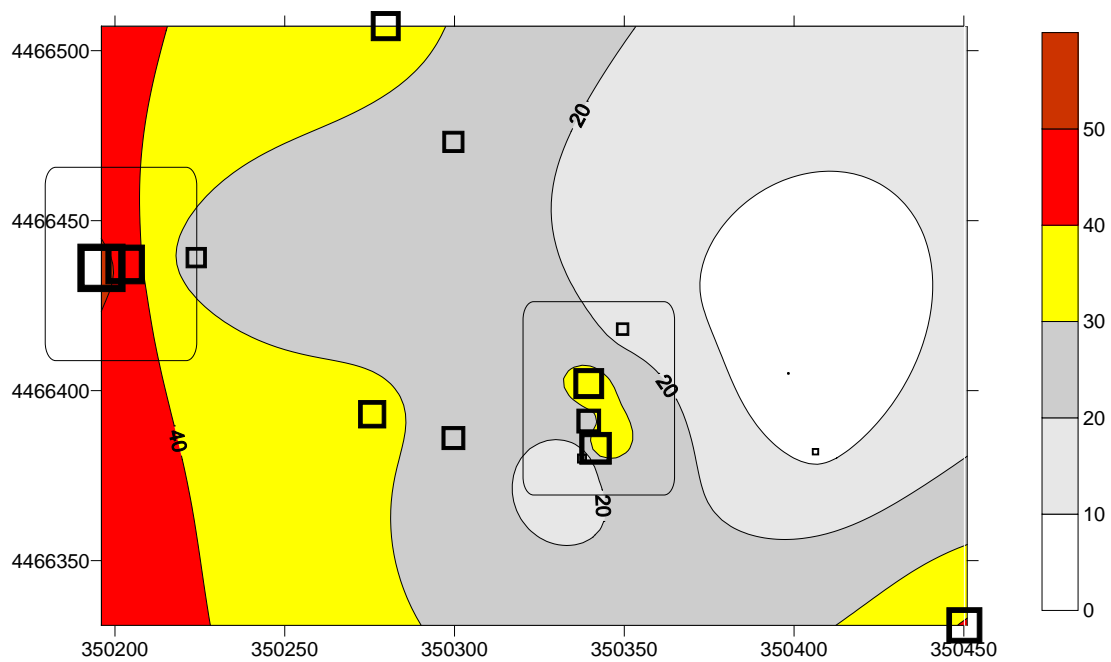


Fig. S5. Growth reduction is spatially associated to locations with fire-scarred trees. The plot shows tree sampled at site CE and the relative growth reduction (contour plot, color scale) expressed as negative growth changes (%). The symbols are trees and their size is proportional to the growth reduction. The two boxes highlight locations where all trees presented conspicuous fire scars (“catfaces”). The axes x (longitude) and y (latitude) correspond to UTM coordinates.

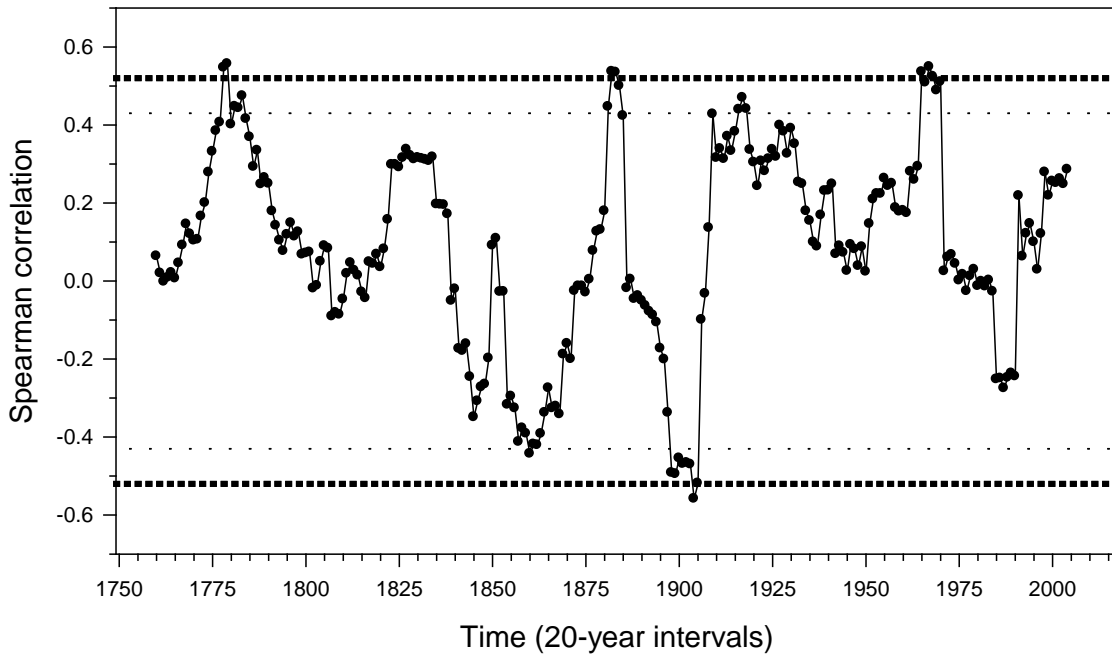


Fig. S6. Spearman moving correlations calculated by relating the annual number of fires according to historical records (Fig. 3a) and short-term changes in growth (differences between standard and residual ring-width indices, see Fig. 4a). Dashed and dotted horizontal lines show the 0.05 and 0.01 significance levels, respectively. Correlations were calculated for 20-year long intervals lagged by one year considering the common 1750-2013 period.