# CA-IDTIMS and biostratigraphy: Their impact on exploration

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## SUMMARY

Uranium-Lead dating of Zircon using the Chemical Abrasion-Isotope Dilution Thermal Ionisation Mass Spectrometry (CA-IDTIMS) technique has largely overcome the problem of radiogenic lead loss, and has greatly improved the precision of the dating. That improvement in precision and the plethora of tuffs throughout the Permian and Triassic successions in eastern Australian basins (Tasmania, Sydney, Gunnedah, Bowen and Galilee) permits the recalibration of the Australian spore and pollen palynostratigraphic scheme directly to the numerical timescale, obviating the need for a multi-step process of correlation. The Permian recalibration has, for example, shown the base of the *Dulhuntyispora parvithola* (APP5) Zone to be about 6 million years younger than previously calibrated, and in the Triassic, the base of the *Polycingulatisporites crenulatus* (APT5) Zone could be about as much as 10 million years younger than previously calibrated.

Key words: CA-IDTIMS, Permian, Triassic, palynostratigraphy, geochronology.

## **INTRODUCTION**

Because of the huge coal reserves present, the Sydney, Gunnedah and Bowen basins have been intensively studied (Figure 1). However, because much of the stratigraphic succession was deposited in non-marine environments, stratigraphic correlation is difficult even within a basin, and the most useful biostratigraphic scheme is based on largely endemic spores and pollen.

Until the development of Chemical Abrasion-Isotope Dilution Thermal Ionisation Mass Spectrometry (CA-IDTIMS), one of the major problems with Uranium-Lead dating of zircons was that of radiogenic lead loss, which gave ages significantly younger than that of crystallisation. Several techniques were developed to try to overcome this problem, including physical abrasion, mild HF leaching and neither of these have been entirely successful. The 'chemical abrasion' method, developed by Mattinson (2005), entailed annealing, followed by digestion in concentrated and pressurised HF over an extended period, and it seems to have largely overcome the problem of lead loss (Figure 2).

#### METHOD AND RESULTS

Previous techniques gave 95% confidence intervals of 1% or worse, whereas CA-IDTIMS can deliver 95% confidence intervals of 0.1% or better. This has major implications for the correlation of strata. Previous techniques allowed the dating of formations, subgroups or groups, but we can now date individual beds in a succession and provide a much better understanding of the timing of volcanic events and sedimentation rates. However, perhaps the most important facet of this new technique is the ability to date biostratigraphic zones. Previously, zones were calibrated against the numerical timescale often by a three-stage correlation. For instance, in the Permian, eastern Australian palynological zones were correlated with Western Australian palynological zones, on the assumption that they were coeval. Then limited Western Australian condont or ammonoid occurrences were used to correlate to northern hemisphere zonal schemes, which form the basis for the international Geologic Time Scale. Each of these steps added a degree of uncertainty that is rarely, if ever, quantifiable. The result is essentially presented as the best available estimate. Where ash beds are common, and they are very common in the eastern Australian coal basins, the new technique allows robust calibration of biostratigraphic schemes directly to the numerical timescale, thus obviating the need for the three-stage, imprecise correlations previously used.

The recalibration of the Guadalupian and Lopingian (middle and late Permian) have been completed (Laurie et al., 2016), based on 34 CA-IDTIMS dates and demonstrate that (see Figure 2): the *Playfordiaspora crenulata* and *Protohaploxypinus microcorpus* (APP6) zones are most likely Triassic in age rather than Changhsingian (i.e. their base being about 2.5 million years younger than previously calibrated); the base of the *Dulhuntyispora parvithola* (APP5) Zone is about 6 million years younger; the base of the *Dulhuntyispora dulhuntyi* (APP4.3) about 7 million years younger; the base of the *Didecitriletes ericianus* (APP4.2) Zone is about 2 million years younger; the base of the *Dulhuntyispora granulata* (APP4.1) Zone is about 1.8 million years younger; and the *Microbaculispora villosa* (APP3.3) Zone is about 3 million years younger. Preliminary results from the Cisuralian (early Permian) are also available (Bodorkos et al., 2016) and indicate that: the base of the *Praecolpatites sinuosus* (APP3.2) Zone is about 2.6

million years younger; the base of the *Phaselisporites cicatricosus* (APP3.1) is about 3.5 million years younger; the base of the *Microbaculispora trisina* (APP2.2) Zone is at about the same age; the base of the *Pseudoreticulatispora pseudoreticulata* (APP2.1) Zone is about 2.3 million years older; and the base of the *Pseudoreticulatispora confluens* (APP1.22) Zone is about 1.4 million years older.

Currently, only seven CA-IDTIMS dates have been determined for Triassic biozones. Of these, five are from eastern Australia, and two from New Zealand. A very preliminary recalibration of eastern Australian Triassic biozones based on these few preliminary dates is given in Figure 3. While most of the new dates are confidently assigned to palynozones (Bomfleur et al., 2014; Smith et al., 2015), one sample, from the Brisbane Tuff of the Clarence-Moreton Basin (Queensland) is not based on direct palynostratigraphic dating, but on a correlation of the Brisbane Tuff with the base of the Ipswich Coal Measures (de Jersey & Hamilton, 1965). It provides an age of  $227.08 \pm 0.10$  Ma for the *C. rotundus* biozone. This suggests that the base of APT4 may need to be recalibrated to approximately 10 million years younger than previously.

Such dramatic changes in the calibration will, of course, have a significant effect on age-depth curves and the resultant geohistory analyses.

### CONCLUSIONS

CA-IDTIMS is a more precise technique for Uranium-Lead dating of Zircons. The eastern Australian Permian and Triassic successions contains numerous tuffs throughout. This permits a more precise and accurate calibration of the largely endemic sporepollen zonation, than has hitherto been possible. It has demonstrated that the calibration obtained from the previous method of multistep correlations needs revision.

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Figure 1. Permian-Triassic Basins in Eastern Australia. Most of the CA-IDTIMS dates used herein have been obtained from the Tasmania, Sydney, Gunnedah, Bowen and Galilee basins (Figure modified from Metcalfe et al., 2015).

Ма	Period	Epoch	Stage	by a (C	Calibration ntle et al. (2010) ted to GTS 2012 stein et al. 2012)	Calibration herein				
251	Triassic	Early	Olenekian _/	APT1		Protohaploxypinus samoilovichii Lunatisporites	APP6		P. microcorpus	
252 <b>-</b>		Ľ	Changhaingian	ian APP6		Protohaploxypinus microcorpus			<b>1</b>	
254 -		Lopingia				Playfordiaspora crenulata		5	D. parvithola	
255			Wuchiapingian				Arr3		M. evansii Acme	
257-				APP5		M. evansii Acme			<b>*</b> 3576	
258-						Dulhuntvispora		4.3	D. dulhuntyi	
259						parvithola	APP4	4.2		
261-		Guadalupian	Capitanian							
262 -									🗕 D. ericianus	
263						Dulhuntvispora			-	
265					4.3	dulhuntyi			*	
266 -			Wordian Roadian	APP4	4.2	Didecitrilates			266.8	
267						ericianus			D granulata	
269-								4.1		
270-						Dulhuntyispora granulata		3.3	M. villosa	
271	C								271.4	
273-	al	uralian	Kungurian	APP3	3.3	Microbaculispora	APP3	3.2	P. sinuosus	
274 -	J.					Viilosa				
275	E.				3.2				276.0	
277-	e					Praecolpatites sinuosus		3.1		
278-	Д								P. cicatricosus	
279					3.1				280.0	
281-			Artinskian			Phaselisporites cicatricosus		2.2		
282-									<b>±</b>	
283-					2.2				M. trisina	
285-				APP2		Microbaculispora				
286-						trisina			297.0	
287-		<u>is</u>			2.1	Pseudoreticulatispora pseudoreticulata	APP2	2.1	207.0	
289-		0								
290-									← P. pseudoreticulata	
291-			Sakmarian							
293-										
294 -					1.22	Pseudoreticulatispora			-	
295-			Asselian			contiuens			295.5	
297				APP1	1.21	Microbaculispora	APP1	1.22	P. confluens	
298-						tentula		1.21	M. tentula	
299 -	Car	Carboniferous			1.1	Protohaploxypinus spp.				

Figure 2. Recalibration of the Australian Permian palynostratigraphic scheme from that of Mantle et al. (2010), based on Laurie et al. (2016) for the Guadalupian and Lopingian, and unpublished data for the Cisuralian.

Ma Period Epoch Age/Stage					Previous Eastern Australian Spore-Pollen Zonation (Mantle et al. 2010)			Preliminary recalibration (Bomfleur et al. 2014 [APT5], Smith et al. 2015, Laurie et al. 2016 and herein)		
201	Jurassic	Early	Hettangian	APJ1		8	APJ1			
207   203   204   205   206   207   208   208			Rhaetian		5.2		APT5		← (NZ) F. moretonensis	
210									P. crenulatus	
211       212       213       214       215       216       217       218       219       220       221       222       223       224	U	Late	Norian	APT5	5.1	P. crenulatus	4.2 APT4		<b>≠</b> C. rotundus	
224 <b>–</b> 225 <b>–</b> 226 <b>–</b>	SSİ			APT4	4.2	C. rotundus				
227	ä								+	
229 230 231 232 233 234 235 236 237	Tria		Carnian		4.1			3.4		
238					3.4		APT3		A. parvispinosus	
239   240   241		Middle	Ladinian		3.3	A. parvispinosus		3.2		
242   243   244   245   246			Anisian	APT3	3.2			3.1		
247 <b>–</b> 248 <b>–</b> 249 <b>–</b>		rly	Olenekian	APT2	2.2 2.1	A. tenuispinosus P. samoilovichii	APT2	2.2 2.1	A. tenuispinosus	
250 <b>—</b> 251 <b>—</b>		Eai	Induan	APT <sup>.</sup> Induan		APT1 L. pellucidus		1	L. pellucidus	
252	Permian	Lopingian	Changhsingian	APP6 APP5		<ul> <li>P. crenulata</li> <li>D. parvithola</li> </ul>	APP5		D. parvithola	

Figure 3. Recalibration of the Australian Triassic palynostratigraphic scheme from that of Mantle et al. (2010), based on Bomfleur et al. (2014), Smith et al. (2015), Laurie et al. (2016), and unpublished data.